Annual Report FY 2002

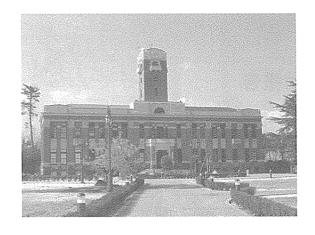
平成 14 年度活動報告

Institute for Geothermal Sciences Graduate School of Science Kyoto University

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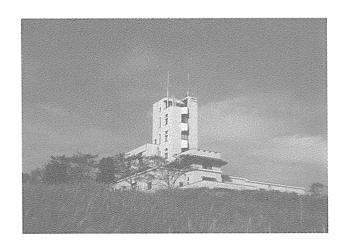
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Front Cover Image:

A strombolian explosion in the 1st crater of Mt. Nakadake, Aso volcano in October 1979. (Photo by M. Sako)

表紙の写真

1979年10月の阿蘇中岳第一火口のストロンボリ噴火の様子(迫幹夫撮影)

Editional compilation by T. Shibata

Printed in Japan by Touka Shobo

目次 Contents

序	Preface		iv
構成員	Members		1
研究活動	Research Activ	rities	
機	関内共同研究	Insider Colaboration	2
研	究報告	Scientific Reports	4
公	表論文	Publications	34
学	会発表	Presentations	38
共同研究		Collaborations	44
定	常観測	Routine Observations	45
装置・設備	fi Instruments a	and Facilities	47
研究費	Funding		49
教育活動	Eductation		50
学	位、授業	Academics	50
セ	ミナー	Seminars	51
学会活動	Activities in	Scientific Societies	53
社会活動	Public Relat	tions	53
_	一般見学会報告	Reports on open house	55
来訪者	Visitors		58

平成 14 年度は大学の法人化への準備と 21 世紀COEプログラムの申請に多大の労力が割かれました。この作業に 1997 年度から出版されている年報が大いに役立ったことを述べて、出版に努力していただいた方々に謝意を表します。

法人化に伴い、学部附属の施設は省令から外れ、大学の裁量で施設経費および存続さえもが決定されることになります。理学研究科で準備した中期目標・中期計画の中で、「研究施設やセンターは、教育・研究において重要な役割を果たしており、さらに充実させる」と明記されてはいますが抽象的であります。地球惑星科学専攻において研究施設の存在をアピールするとともに、役割についても理解を深めておくことが必要です。21世紀COEに関しては、地球熱学研究施設は地球惑星科学専攻とともに、「活地球圏の変動解明:アジア・オセアニアから世界への発信」を申請し、主に「マグマ活動と水・熱フロー」のサブタイトルで貢献しようとしていますが、多目的観測ネットワークや、活地球圏情報統合ポータルの立ち上げ等にも積極的な協力が要請されています。これらの作業の中で、理学研究科からは業績批判、学生不足が指摘され、一層の整理統合を望む発言もありました。批判は謙虚に受け止める必要があり、施設構成員各自が学問的視野に立って研究施設の将来像を構築してゆく必要があります。

火山関係では第7次火山噴火予知計画が策定中であります。火山研究センターは、地元はじめ 火山研究者達からの絶大な信頼を得て阿蘇火山の観測に携わっており、カルデラ内にボアホール 地震計を設置するなど観測設備の充実を企画しています。阿蘇火山は2000年秋以降、中岳火口の 一部が赤熱現象を呈しています。火山活動把握と噴火機構の解明を進展させる良い機会になるで しょう。

教育面では、学生および教室との交流を活発にするために、地球熱学、火山物理学の毎週講義を開始するとともに、野外実習として別府における室内実習を加えました。より多くの学生諸君が興味を抱き、熱学に関わる研究を指向してくれることを期待しています。

人事面では、外国人客員のAnthony Hurst 氏が9月末でニュージーランドに帰り、陳中華氏が、研究課題「北部琉球火山フロントにおけるマグマの生成」で10月に着任されました。また、同位体分析の中心であった鈴木勝彦助手がJAMSTEC に転出され、非常勤研究職員では宮田佳樹、宮崎隆両氏が去りました。 阿蘇では橋本武志助手が2月末で北海道大学助教授として転出され、宇津木充氏が4月から助手として着任されました。さらに、3月末には小野博尉助手が定年退官されました。同氏は阿蘇火山の火口周辺地震観測網の整備や、山上地震データの無線テレメータ化に尽力され、火山性微動や発震機構の研究を進展させるとともに、阿蘇火山や雲仙火山の地殻変動観測にも貢献されました。長年のご尽力に感謝いたします。

この活動報告が施設全員の相互理解と研鑽の拠り所として活用されることを願い、施設外の方々には、私たちの活動をご理解戴き、ご指導ご助言を賜る機会となれば幸いです。

平成 15 年 5 月 平成 14 年度地球熱学研究施設長 田中良和

Preface

In the financial year of 2002, we spent many time to make proposal for the 21st century COE and preparation against reform to juridical personalization of national universities. It is my pleasure to write the annual reports published since 1997 were very useful in these works.

After the reorganization, budget of our laboratory is no warrant and continuance is free from a ministerial ordinance. In the middle period plan arranged by the faculty of science, necessity of more complete support was promoted as the laboratory functions well in education and research, however, it is too abstractive. It is important to appeal our science study, and discuss about significance of our laboratory with the members in Kyoto campus.

For COE, our group joined to the major of earth and planet science with the project "Study of fluctuations in dynamical earth: Call to the world from Asia and Oceania", and our contribution is requested in a subject of " relations between magma activity and water/ heat flow ", also promotion of observational network and potal of information system for active earth sciences are entreated. In the process, the top-level of the faculty criticized few students and lack of publications and claimed to make new project with the staff in main campus.

We must accept severe criticism. Each staffs should consider strategy from the scientific point of view, and create a future plan through active discussion.

The 7th national program for prediction of volcanic eruption is drawn up now. As AVL get reliance on volcano monitoring of Aso by residents and researchers, we apply new equipments such as borehole seismometer. Status of Aso volcano keeps in red heat at the southern end of 1st crater since autumn in 2000. It gives us a good chance to investigate such mechanisms through detailed observations.

In education, weekly lectures of geotherm and volcanology were started aiming active good wishes with students and staffs in Kyoto, moreover practical training course of material analysis was appended at Beppu.

As personnel reshuffle, Dr. Anthony Hurst replaced by Dr. Chang-Hwa Chen, who works on magma generation in northern Ryukyu volcanic front. Associate Dr. Katsuhiko Suzuki, one of the most key person in isotope geochemistry, moved to JAMSTEC, and Dr. Takeshi Hashimoto was promoted to associate professor of Hokkaido University. Associate Hiroyasu Ono was retired in March. We appreciate his many contributions making up telemeter systems around the crater of Aso, research works of tremor and seismic mechanism and observations in ground deformation on Aso and Unzen. In the case of the part time instructors, Dr. Yoshiki Miyata and Dr. Takashi Miyazaki resigned on 31 March.

I wish this annual report be put to practical use for improving the mutual understanding among the staff and providing study material for our daily activities. I would like to entreat all people to read it to understand our activities and we look for guidance on further development of the Institute.

> Aso, May 2003 Yoshikazu Tanaka, Professor/ Director



構成員 Members

教授	Professors	Anthony Hurst					
竹村恵二 Keiji Takemura		2002 年 9 月 退職					
田中良和	Yoshikazu Tanaka		a				
(施設	長 Director)	非常勤研究員	Resarch Associates				
由佐悠紀	Yuki Yusa	井上(北田)	直人 Naoto Inoue (Kitada)				
		西村光史	Koshi Nishimura				
助教授	Associate Professors	宮田佳樹	Yoshiki Miyata				
大沢信二	Shinji Ohsawa	宮崎隆	Takashi Miyazaki				
大倉敬宏	Takahiro Ohkura						
須藤靖明	Yasuaki Sudo	教務補佐員	Research Assistant				
古川善紹	Yoshitsugu Furukawa	芳川雅子	Masako Yoshikawa				
助手	Assistant Professors	研究支援推進員 Technical Assistants					
宇津木充	Mitsuru Utsugi	増田秀春	Hidearu Msuda				
20	03 年着任						
小野博尉	Hiroyasu Ono	大学院生	Graduate students				
川本竜彦	Tatsuhiko Kawamoto	網田和宏	Kzuhiro Amita				
柴田知之	Tomoyuki Shibata	長谷英彰	Hideaki Hase				
鈴木勝彦	Katsuhiko Suzuki	中坊真	Makoto Nakaboh				
20	03 年退職	山田 誠	Makoto Yamada				
橋本武志	Takeshi Hashimoto	吉川美由紀	Miyuki Yoshikawa				
20	03 年転出	新名真裕美	Mayumi Shinmyo				
技官	Technical Professionals						
迫 幹夫	Mikio Sako	事務補佐員	Secretaries				
馬渡秀夫	Hideo Mawatari	今村町子	Matiko Imamura				
吉川慎	Shin Yoshikawa	後藤君子	Kimiko Goto				
		土屋寿子	Toshiko Tsutiya				
外国人客員	Visiting Faculty						
陳 中華		臨時用務員					
20	002年10月着任	山崎咲代	Sakiyo Yamazaki				

研究活動 Research Activities

機関内共同研究 Insider Colaboration

Seismic activity in the Beppu graben, Kyushu, Japan.

T. Ohkura, Y. Furukawa, K. Takemura and H. Mawatari

The Beppu graben is located at the eastern end of the Beppu-Shimabara rift zone in Kyusyu, Japan, where E-W striking normal faults are dominated. There are three active volcanoes in the graben, named Mt. Yufu, Mt. Garan and Mt. Tsurumi (Fig. 1).

Since March 1993, seismic activities have been monitored by BGRL, Kyoto University with five seismic stations in and around the graben (Furukawa et al., 2000). Preliminary hypocenter distribution was obtained by Furukawa et al. (2000) using P-wave arrival times. However, in order to determine hypocenters with more accuracy, S-wave arrival times should be used. In this study, we relocated hypocenters using P and S arrival times in order to discuss relation between seismic activities and tectonics in this region.

We could relocate more than 500 earthquakes, which occurred from October 1993 to December 1999, recorded at four or five stations. In Fig. 2, epicenter distributions with E-W and N-S cross sections are shown. In this figure, hypocenters located with the RMS residual time less than 0.15

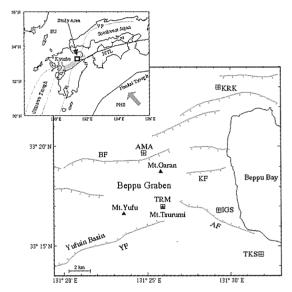


Fig. 1 Tectonic map of the study area. The rectangles and the solid triangles denote the seismic stations and the active volcanoes, respectively. The shaded lines indicate active faults. BF: Beppu-kita Fault. KF: Kannawa Fault. YF: Yufuin Fault. AF: AsamigawaFault

second are plotted. We can recognize that almost all the events are located in the graben which is bounded by the Beppu-kita fault at the north and the Yufu –Asamigawa faults at the south as shown in Fig. 2.

As shown in the E-W cross section of Fig. 2, the lower limit of seismicity becomes shallower toward the eastern flank of the active volcanoes; Mt. Garan and Mt. Tsurumi and there is an aseismic zone at a depth more than 5 km. A high electric conductive body revealed by electromagnetic survey (NEDO, 1989) is located in the upper part of the aseismic zone. From these features, this aseismic zone could be considered to be a high temperature body.

As shown in Fig. 3A, many earthquakes with a depth less than 3km are located in the region between the Kannawa fault and the Asamigawa fault. In this region, there are many hot springs and hydrothermal circulation of meteoritic water is very active. The effects of fluids on rock strength possibly

cause these earthquakes.

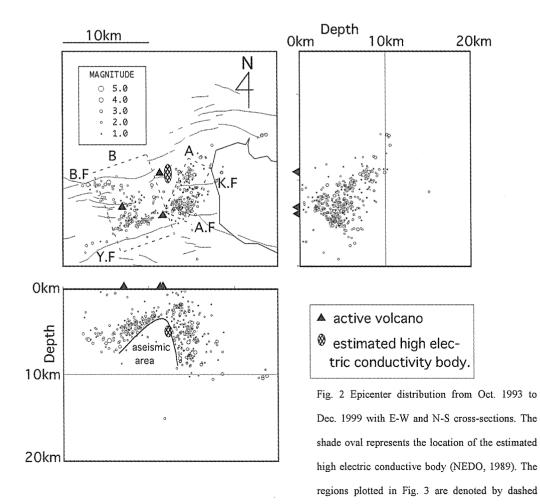
Epicenters can be delineated along the Yufuin Fault about 2km north to the surface trace of this fault as shown in Fig.3B. Taking into the depth of these events, a large dip angle of this fault can be estimated. Although some earthquakes show the first arrival polarities which are consistent with right lateral movement along this fault (Fig. 3B), mechanism solutions in this region could not be determined uniquely by the five stations data. Therefore, further studies on focal mechanisms combining the data from adjacent seismic network are necessary for a better understanding on tectonics and structure of this fault that is the southern boundary of the Beppu-Shimabara graben.

Acknowledgments

HyperDPRImap ver. 3.9.1 by Dr. H.P. Katao was used to draw Figs. 2 and 3.

References

Y. Furukawa, H. Mawatari and T. Kuroda (2000) Seismic Activities in Beppu, Kyushu, Zisin 2, 53, 487-490
 NEDO (1989) Regional exploration of geothermal fluid circulation system, Tsurumi-dake area. National Geothermal Resources Exploration Project, New Energy Development Organization, Tokyo. pp342



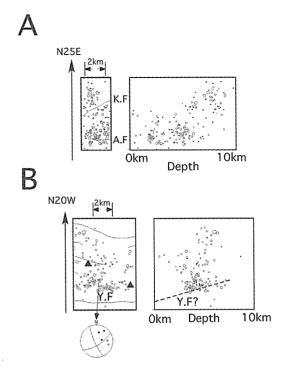


Fig. 3

- A: Epicenter distribution in the region A in Fig.2 and vertical cross section along N25E, almost perpendicular to the Yufuin fault (Y.F).
- B: Epicenter distribution in the region B in Fig.2 and vertical cross section along N20W, almost perpendicular to the Asamigawa fault (A.F).

研究報告 Scientific Reports

Mixing Process of Air and Underground Water into Magmatic Gas Discharged from Kuju-Iwoyama Fumarolic Area of Kuju Volcano, Central Kyushu, Japan

K. Amita and S. Ohsawa

Steams and gases discharged from fumaroles at discharge temperature ranging from 98°C to 308°C in Kuju-Iwoyama fumarolic area of Kuju Volcano were sampled without atmospheric air contamination for isotopic and chemical analyses. Stable isotope ratios of hydrogen and oxygen (δ D and δ ¹⁸O) of the steam condensate samples showed that the fumarolic steams were mixtures of magmatic steam and local meteoric water. Moreover, the He-N₂-Ar contents of the fumarolic gas samples indicated that they consisted of magmatic gas, air and air-saturated water (ASW). From the quantitative estimation by means of the δ D and δ ¹⁸O values and He/Ar ratios of the fumarolic gases, the proportion of mixing for the three end-members: magmatic fluid_(g), air_(g), ASW₍₁₎ was estimated to be 1 : 0.0005-0.0021 : 0.0009-0.0043. This relative proportion and a linear relationship between the δ -values (δ D and δ ¹⁸O) of steam and He/Ar ratios which are corrected to be free from ASW contribution strongly suggest that the underground air and the underground water can be treated as the vadose air and the vadose water, respectively. From above discussions they suggested that the fumarolic fluids of Kuju-Iwoyama would be formed by the mixing of the vadose fluids (vadose air and vadose water) and the magmatic fluid into the fumarolic gas conduits within the vadose zone. This mixing model (Fig.1) will reasonably explain the

contradiction in terms of the isotopic and chemical data of fumarolic steams and gases which are obtained from several active volcanoes (e.g. Satsuma-Iwojima Volcano, Kuju-Iwoyama between 1959 and 1984); δ D and δ ¹⁸O of steam condensates show the typical values of andesitic magmatic steam, whereas He/Ar ratios of fumarolic gases clearly indicate the mixing of air and/or ASW into magmatic fluid.

(Submitted to J. Geotherm. Res. Soc. Jpn.)

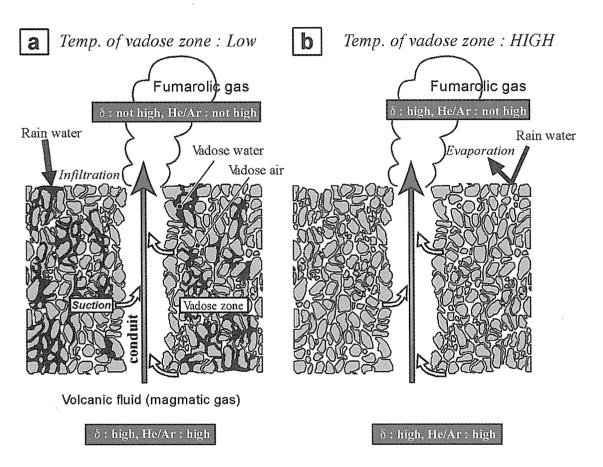


Fig. 1 (a) Schematic cross-section depicting mixing process between vadose fluids (water + air) and volcanic fluid (magmatic gas) rising through fumarolic gas conduit, with information on change in δ-value (δ D and δ ¹⁸O) and He/Ar ratio before and after mixing of vadose and volcanic fluids.
(b) Schematic explanation on change in δ-value and He/Ar ratio before and after mixing of vadose and volcanic fluids when temperature of vadose zone is high.

Precise determination of rare earth elements and yttrium in GSJ standard rock sample JP-1 (Peridotite) by microconcentric desolvating nebulisation ICP-MS

O. Chang, T. Shibata, K. Shinotsuka, M. Yoshikawa and Y. Tatsumi

Compiled reference values of trace and rare earth element (REE) concentrations for GSJ standard rocks JP-1 (peridotie) are mostly preferable rather than recommended ones. Major limitations for precise and accuracy determination of these elements for such mantle-derived material as JP-1 are its extremely low concentration (at ng g⁻¹ level) and ultramafic properties. We estimated the REE and Y results of JP-1 obtained using ICP-MS coupled with a microconcentric desolvating nebulizer.

Digestion of powder sample followed a modified HF/HClO₄ method. 100 mg powder was first treated with HF/HNO₃ under ultrasonic bath, then attacked by HF/HClO₄ on hotplate, fumed 1~2 times with HClO₄. Residuals was re-dissolved with 2% HNO₃ and finally spiked with ¹¹⁵In and ²⁰⁹Bi for internal standardization before ICP-MS measurement. The whole procedure was done in a class 100 clean room. Consequently, blank contributions were reduced to less than 1% for all analyte except Eu and Sm which were around 3% and are within analytical uncertainty. Such low blank level is fatal important when ultra low abundance samples are measured.

A low flow desolvating nebulizer (Aridus[®] MCN6000; CETAC) was used for sample introduction to ICP-MS. Optimization of the instrument was done by usual tuning balanced with sweep Ar gas and N₂ gas flows to MCN6000 unit. The sensitivity achieved were 430000 and 270000 cps/ppb for ¹¹⁵In and ²⁰⁹Bi, respectively, which are about 10 times higher than equivalent instrument in normal operating mode.

Moreover, oxide levels are reduced by 2 orders, so that middle REE are avoided from oxide interferences, such as BaO on Eu and PrO on Gd.

Our new data for JP-1 are presented in Table 1. It is evident that they produced a smooth, coherent chondrite normalized REE pattern (Figure 1). Our LREE and HREE values are generally consistent with those reported by Makishima and Nakamura (1997), but MREE from Sm to Ho are lower than theirs. However, the new data for JP-1 are critically different to

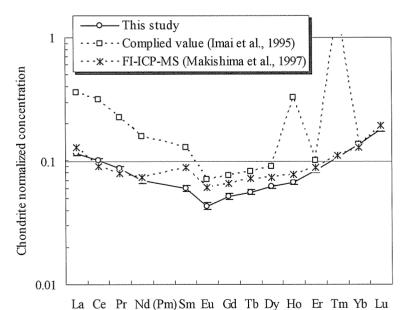


Figure 1. Measured REE for JP-1 by ICP-MS coupled with CETAC Aridus® MCN 6000. Error bar shown for each data point is 1s of 3 individual analyses.

suggested values of GSJ (Imai et al., 1995). REE pattern obtained here does not exhibit convex curve of

LREE, nor anomalies of Ce and Sm. With the high sensitivity and desolvating performance of ICP-MS fitted with MCN6000, the internal precisions were significantly improved, especially for MREE which are 3-8% (1s). These MREE precision levels are equivalent to or better than those obtained by sector high-resolution ICP-MS for same type samples (PCC-1 and DTS-1; Jain et al., 2000). The overall reproducibility with this procedure are 2.2-4.8% (1s) for REE except for Nd, Sm, Eu and Gd which are 5.5-6.6% (1s). On the basis of demonstrated level of precision, accuracy and the REE pattern, we conclude that these new values for JP-1, obtained by ICP-MS coupled with MCN6000 are the best estimates.

Sample: JP-1		Y	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Но	Er	Tm	Yb	Lu
digestion 1	ppm	0.096	0.026	0.057	0.0073	0.029	0.0085	0.0025	0.010	0.0021	0.015	0.0038	0.013	0.003	0.022	0.0044
**	rsd, %	3.1	2.6	1.6	1.4	3.7	3.9	7,6	7.4	4.5	4.7	3.6	4.9	4.2	3.8	2.4
digestion 2	ppm	0.091	0.028	0.063	0.0079	0.032	0.0089	0.0025	0.010	0.0020	0.015	0.0037	0.013	0.003	0.022	0.0043
	rsd, %	1.2	0.8	1.7	2.0	3.5	8.7	4.1	5.9	2.0	2.5	1.8	3.4	3.1	2.0	1.2
digestion 3	ppm	0.091	0.028	0.061	0.0079	0.032	0.0090	0.0023	0.010	0,0020	0.015	0.0037	0.014	0.003	0.022	0.0043
-	rsd, %	1.0	1.2	0.3	1.8	3.3	2.3	2.3	4.0	5.2	1.7	3.6	3.9	3.5	1.0	2.2
Mean, JP-1	ppm	0.093	0.027	0.060	0.0077	0.031	0.0088	0.0024	0.010	0.0020	0.015	0.0037	0.013	0.003	0.022	0.0043
	RSD, %	3.1	4.5	4.8	4.1	5.5	5,9	6.6	5.7	4.5	3.4	3.8	3.9	3.9	2.6	2.2
Makishima and																
Nakamura, 1997	ppm	0.100	0.030	0.054	0.007	0.033	0.013	0.003	0.013	0.003	0.018	0.004	0.014	0.003	0.021	0.005
lmai et al, 1995	ppm	1.540	0.084	0.190	0.020	0.072	0.019	0.004	0.015	0.003	0.022	0.018	0.016	0.041	0.022	0.004
Total precedural bla	ınk, pg	33	27	32	6	24	28	10	17	1	13	3	8	3	22	6
Blank contribution.	%	0.3	0.8	0.4	0.7	0.6	2.7	3.4	1.4	0.5	0.7	0.6	0.5	0.8	0.8	1.2

Laboratory experiments of zeta potential for rocks from some volcanoes in Kyushu district

H.Hase, T.Ishido*, T.Hashimoto, W.Kanda**, S.Sato*** and Y.Tanaka
*Institute for Geo-Resources and Environment, Geological Survey of Japan, AIST
**DPRI, Kyoto Univ.

***Instutue for Frontier Reseach on Earth Evolution, JAMSTEC

Self-potential (SP) method has been widely used in volcanic and geothermal areas since it was successfully applied to Kilauea by Zablocki et al. (1976). One of the physical basis of these studies is the streaming potential that arises in fluid flow in a porous medium. In a microscopic view, the electrical double layer of a solid-fluid interface and relative motion of fluid to the solid phase plays fundamental role for generating the streaming potential. ζ -potential (potential of the slipping plane) characterizes this electrical double layer and is one of the most important parameters that control the macroscopic manifestation of SP on the ground surface. It is hence important to know the ζ -potential of rocks in a target field to evaluate any SP distribution either in qualitative or quantitative manner. It is commonly known that the ζ -potential usually takes a negative value for Si-rich materials such as granite (e.g. Ishido and Mizutani, 1981). As a matter of fact, Si is the major content of general crustal rocks. ζ -potential of the real field, therefore, in many cases, is practically assumed to be negative according to such laboratory experiments for the typical samples. However, Hase et al. (2002) revealed that some rocks from Aso volcano, in contrast, have positive values of ζ -potential. In the present study, we extended our measurements to the volcanic

rocks from some other volcanoes in Kyushu island to clarify how common the positive ζ -potential is in volcanic fields.

We sampled the specimen form four volcanoes where SP maps have been already obtained in the past. We took 48 rock samples for the ζ -potential measurements from Kuju (12), Kaimon-dake (11), Unzen (7), and Sarubayama (8). The experimental conditions are the same as those of Hase et al. (2002), where temperature, electrolyte concentration and pH range are 30°C, 10^{-3} mol/l KCl, and 4 to 10, respectively. As a result of the experiments, positive ζ -potentials were obtained in most samples from Kuju and some specimen from Unzen and Sarubayama. We also measured the chemical content of major elements of these samples to see the relationship to the ζ -potential. We found that the ζ -potential is roughly controlled by the content of rock elements, which supports our previous result from Aso volcanics. Summarizing the present study, the positive ζ -potential is not rare in volcanic rocks, and hence, we need to be careful of it when interpreting the SP map.

On the annual variations in geomagnetic differences observed in Long Valley Caldera, California

Hashimoto, T., M. J. S. Johnston, Tanaka, Y., Sasai, Y., Utsugi, M. and Sakanaka, S.

We deployed eleven proton magnetometers in Long Valley Caldera in order to detect anomalous geomagnetic changes of tectonic or volcanic origin during the period of 1999-2001. Total geomagnetic field differences observed in LVC show clear annual changes up to a few nano-Teslas in amplitude. These can be related to the atmospheric temperature-driven variations in local magnetization. It has been revealed by our investigation that magnetization change due to the shallow ground temperature is the most plausible cause for the annual change. We corrected these annual changes with a simple 1-D thermal diffusion model. Further noise reduction of about a factor of five was achieved with predictive filtering using three component geomagnetic data. The resulting data indicate it is unlikely that there was any anomalous geomagnetic field change from triggered deformation in Long Valley related to the distant Hector Mine earthquake (Oct. 16, 1999). However, there still remains a possibility that changes in long-term trends represent anomalous changes due to the volcanic activity of the area.

Based on the knowledge of the present study we propose that improved understanding of volcanomagnetic or seismomagnetic effects will result from measurements of ground temperature together with the geomagnetic total field. In order to improve precision of repeated geomagnetic surveys by correcting for the effects of annual changes of thermal origin, we should conduct continuous geomagnetic and temperature observations for at least one year as to estimate the relationship (relative amplitude and phase lag) between the two. Once the parameters are obtained, the continuous recording of geomagnetic field is not necessarily required just repeated measurements of magnetic field with the ground temperature. We can then make temperature corrections to the magnetic data. This procedure will considerably reduce the cost and effort to keep a continuous magnetic station because the temperature measurement is generally inexpensive and easy to conduct compared to continuous magnetic recordings.

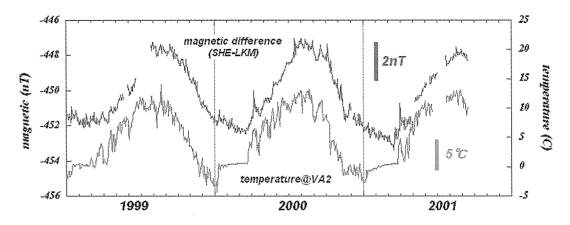


Fig. 1: Daily mean geomagnetic difference of the total force between SHE and LKM and ground temperature plots since Jan. 1991 to Sep. 2001.

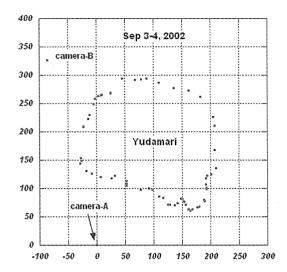
Water level of the crater lake (Yudamari) of Nakadake, Aso Volcano

Hashimoto, T., Ikebe, S., and Yoshikawa, S.

The 1st crater of Nakadake, Aso volcano has a hot crater lake called *Yudamari*. Such a crater lake with a high temperature, discharging considerable amount of thermal energy in terms of vaporization, is a huge cooling wheel in the heat transport system of the volcano. Historical records tell us that *Yudamari* tends to be dried up before the volcanic activity reaches to its climax with a Strombolian eruption. Monitoring the geophysical parameters such as water level, temperature and so on is, hence, one of the effective ways of predicting the volcanic activity in the near future.

The area of *Yudamari* as well as the surface temperature is necessary to estimate the heat discharge from the lake surface. For this purpose we have hitherto referred to a GSI volcano map (scale 1:5000, based on an aerial photo on Oct. 11, 1981). According to the GSI map, water level and area of *Yudamari* in Oct., 1981 were 1154.7 [m] (a.s.l.) and 2.1×10^4 [m²], respectively. Meanwhile, field measurements by the authors on Sep. 3, 2002 resulted in 1145 [m] and 4.0×10^4 [m²], respectively. It is obvious from these results that the 1st crater has been significantly eroded down in these two decades. Although most part of this erosion of the crater wall (widening of the crater) is probably a by-product of the major eruption during 1989-1990, topography of the crater is still changing. It is hence important from a viewpoint of heat discharge estimation to conduct a regular measurement of crater topography. We estimated the heat discharge from *Yudamari* surface using the formula of Ryan et al. (1974) as 150~200 [MW] in Sep., 2002.

Water level of *Yudamari* is decreasing about 5 [m] per year since 2000 as is shown in Fig. 3. This is much related to the recent temperature increase of the lake water accelerating the vaporization at the lake surface. We consider that these changes are essentially due to the increase of heat flow supplied to *Yudamari*.



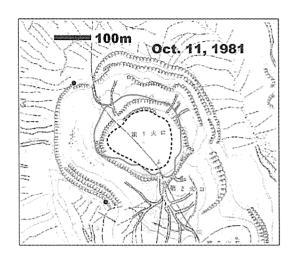


Fig. 1: Edge of the crater lake (Yudamari) on September 3 to 4, 2002 obtained by distance and dip measurements. Units of the axis are in meters.

Fig.2: Topography map (GSI, 1981) of the crater area. Dashed and thick solid lines indicate the edge of Yudamari of Oct. 11, 1981 and Sep. 3-4, 2002, respectively.

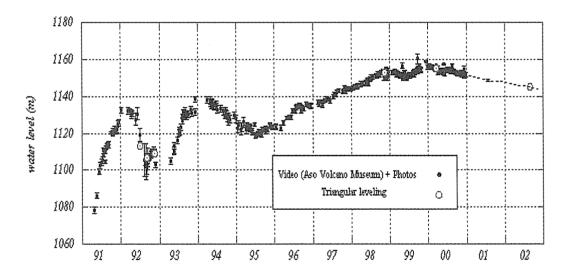


Fig.3: Water level of Yudamari obtained by video image processing and triangulation.

Fault-bounded basement block of Northeastern part of Osaka-wan Fault, Southwest Japan based on Integrated Study of Gravity and Seismic Reflection Data

N. INOUE, N. KITADA and K. TAKEMURA

After the 1995 Hyogo-ken Nambu (Kobe) earthquake, several kinds of geophysical and geological surveys were performed in and around Osaka Bay, Southwest Japan (Fig.1). These surveys, such as seismic reflection, gravity explorations and borehole drilling, revealed the existence of many active and inactive faults in this district. The 40-km long Osaka-wan fault, running from NE to SW, is one of the major active faults in this area. The vertical displacement of this fault is about 1000 m. The survey results suggest that the northeastern part of the Osaka-wan fault splits into three branches in offshore Kobe (Fig. 2). The Osaka Group, Pliocene-Pleistocene soft sediments, covered the basement, which consists of Mesozoic granite rocks. Fifteen marine clay layers in the Osaka Group, from Ma-1 to Ma13 in ascending order, are useful as key beds because they are laterally continuous layers and because their ages are well known by assigned to the Marine Isotope stages. Furthermore, these marine clay layers are well identified on seismic sections due to the contrast in physical properties between marine clay and other sediments.

In this study, marine clay layers obtained from borehole stratigraphy have been correlated to reflectors on seismic sections around the northeastern part of the Osaka-wan fault. The basement configuration was obtained from gravity anomalies with constraints of the seismic reflection data and indicates detailed branched shape (Fig. 3). The depth of the basement exceeds over 2500 m in offshore Kobe and becomes shallower towards Kobe. The estimated sedimentation rate is about 0.8 m/kyr at the deepest basement point in the study area. The sedimentation rate decreases towards Kobe.

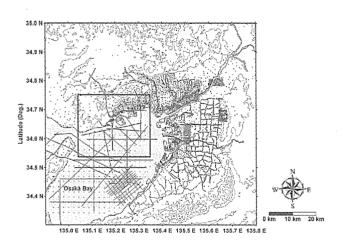


Fig.1 Index map of geological and geophysical survey in Osaka Basin (modified from Inoue and Nakagawa, 2000; Itoh et al., 2001). Black line and solid circle indicate the seismic survey line and gravity station, respectively.

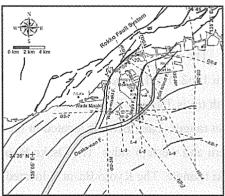


Fig.2 Index map around Kobe. Black thick lines denote the active fault on land (Okada and Togo, 2000) and geological structure in bay (Yokokura et al., 1998; 1999), respectively. Solid circles and lines indicate the location of borehole and seismic survey line used in this study.

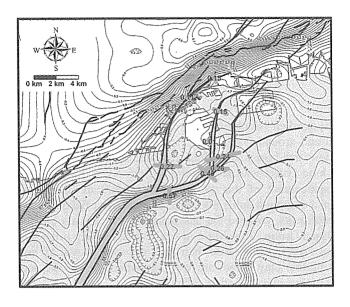


Fig.3 Activity of Osaka-wan fault. Gray zones in figure show the location of seismic section used for estimation of activity. Numerical values on brown lines are activity of fault in m/kyr. Italic numerical values on brown lines are activity of fault estimated by Yokokura et al. (1998). Black thick lines denote the active fault (Okada and Togo, 2000) and geological structure in bay (Yokokura et al., 1998; 1999), respectively. Open circles and rectangles are the gravity station and source of the basement obtained from the seismic profiles, respectively. The numerical value of contour is the altitude of the basement in km. Contour interval is 0.1 km.

Geological Structure in Kinki District Japan, Inferred from Gravity Anomaly

N. INOUE, N. KITADA and K. TAKEMURA

The major neotectonic feature in the Kinki district of central Japan is the Kinki triangle proposed by Huzita in 1962 (Fig.1). The southwest and southeast corners of the triangle are western Osaka Bay and Ise Bay. The north tip is Tsuruga Bay. The Kinki triangle is characterized by the arrangement of small basins and short mountain ranges. Lake Biwa is located in the north of the triangle and is surrounded by many active faults on its eastern and western sides. The Osaka Basin and Nobi Plain are along the northwest and northeast sides of the Kinki triangle. The Kyoto Basin is located between Lake Biwa and the Osaka Basin. Huzita (1974) divided the fault systems, which were developed in the Kinki Triangle and the surrounding areas, into four systems based on their trends and characteristics. The first is a left-lateral fault system with a NW-SE trend, the second is a right-lateral fault system with a dip-slip components along a NE-SW trend, the third is a thrust system with a N-S trend, and the fourth is a system of fracture zones developed along the boundaries of geological units.

AIST (the National Institute of Advanced Industrial Science and Technology, the old name is the Geological Survey of Japan: GSJ) has published the CD-ROM of digital gravity data covering the whole of Japan with a 1km grid in 2000 (GSJ, 2000). According to the gravity grid data, a gravity low appears along the

Osaka Basin and along the shape of Lake Biwa (Fig.2). The gravity low in Lake Biwa is less than -60 mgal and is the lowest anomaly in the Kinki triangle. This gravity low extends to the Chubu mountain ranges and Ise Bay. The thickness of sediments of the Osaka Basin and Nobi Plain are more than 3km and 2km, respectively. The gravity low suggests the presence of the thick sediment. On the contrary, the seismic survey in most parts of Lake Biwa revealed thin sediment of less than 1 km. The gravity low of Lake Biwa suggests the thickening of the crust or subducting of the low-density plate.

In this study, we attempt to detect the shallow and deep structure around the Kinki triangle by regional-residual separation of gravity anomaly. The regional gravity component was derived by separation methods, trend surface analysis and upward continuation. The regional component was obtained based on a deep structural model of the crust and uppermost mantle compiled by Ryoki (1999). In this poster, we discuss the relationship between the shallow gravity component and the subsurface geology and structure, using several analysis methods.

Regional gravity anomaly

The regional gravity anomalies extracted by several methods show similar trend. The high regional gravity anomalies appear in northern Wakasa Bay and southern Kii Peninsula. The low regional gravity anomaly extends from the Inland Sea to the Chubu mountain ranges. The regional gravity anomalies obtained by mathematically show high anomaly extending from Osaka Plain to Kii Peninsula. The location of the lowest regional gravity anomaly based on the model is the Chugoku mountain ranges, on the contrary, mathematical regional anomalies show the lowest anomalies in Chubu mountain ranges. Compared with the regional gravity anomaly obtained by constrained trend-surface analysis (Fig.3), mathematical regional anomaly seems to be better to represent the subsurface structure in the sedimentary basins of Kinki district.

Residual gravity anomaly

The low residual gravity anomaly (Fig.4) extends along the eastern and western sides in the Kinki triangle. The thickness of sediments of the Osaka Basin and Nobi Plain are more than 3km and 2km, respectively. The low residual gravity anomaly suggests the presence of the thick sediment. On the contrary, the seismic survey in most parts of Lake Biwa revealed thin sediment of less than 1 km. The low residual gravity anomaly of Lake Biwa suggests the thickening of the crust or subducting of the low-density plate. The southern side of the Kinki triangle indicates the high residual gravity anomaly. Nara Basin denotes the high residual gravity anomaly. The northern high residual gravity anomaly outside of the Kinki triangle corresponds to the distribution of Paleozoic rocks.

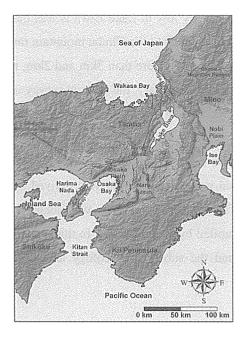


Fig.1 Topography of Kinki district. Topographical data was adopted from 250m-Digital Elevation Model provided by Geographical Society of Japan. Black lines in figure indicate the active fault adopted from Okada and Togo (2000).

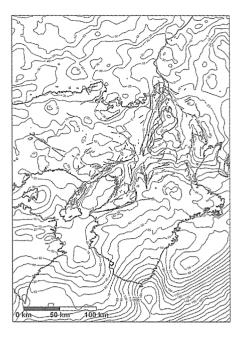


Fig.2 Bouguer anomaly map in Kinki District. Bouguer anomaly data was meshed from 1km mesh of Gravity CD-ROM published by AIST in 2000 with grid interval of 2km. Bouguer correction density is 2.67 g/cm³. Contour interval is 10 mgal. Black lines indicate the active fault by Okada and Togo (2000).

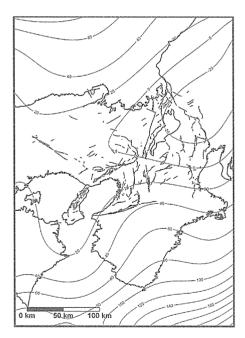


Fig.3 8th order polynomial trend-surface. Contour interval is 20 mgal. Black lines indicate the active fault by Okada and Togo (2000).

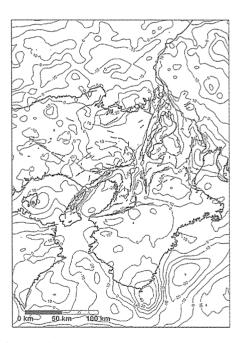


Fig.4 Residual anomaly subtracting 8th order trend-surface (Fig.3). Contour interval is 10 mgal. Black lines indicate the active fault by Okada and Togo (2000).

Subsurface Structure around Unzen Volcano based on Gravity and Geological Data N. Inoue and K. Takemura

The Unzen Scientific Drilling Project (USDP) reveals the subsurface structure, geological history and physical property of Unzen volcano based on acquired various and significant geological, geophysical and geochemical data. We inferred the outline of the subsurface structure around Unzen volcano by 2-D gravity data analysis with constraints of the geological data. The Bouguer anomaly data around Unzen volcano (Fig.1) were adopted from the gravity CD-ROM published by the Geological Survey of Japan (GSJ, 2000).

The two-layered models along with the profiles A to D shown in Fig.1 were constructed. The density contrast is 0.3 g/cm³ (NEDO, 1984). The control points were adopted from Hoshizumi et al. (2002). The subsurface structure was obtained by correcting the model in proportion to the residual gravity. The obtained profiles are shown in Figs. 2 to 5. The A profile (Fig.2) indicates the graben shape of the basement. The B profile (Fig.3) indicates the convex shape of basement under the Mt. Fugen. The C profiles (Fig.4) and D (Fig.5) indicate the similar convex basement configuration under the Mt. Fugen. The obtained basement structures are accordance to the top surface of the basement configuration of the Hoshizumi et al. (2002) except for the profile B.

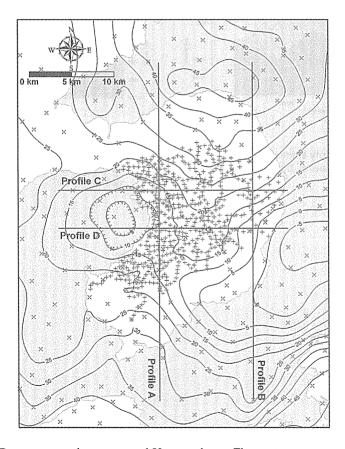


Fig. 1 Bouguer anomaly map around Unzen volcano. The contour map was created from gravity data published by GSJ (2000). The Bouguer correction density is 2.30 g/cm 3 (NEDO, 1984). Plus and cross marks indicate the organization of original gravity data. +: NEDO. $\times:$ GSJ.

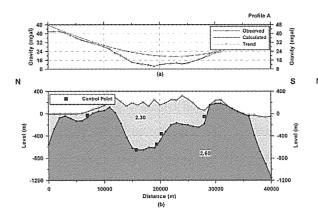


Fig.1 Profile A (a) Gravity anomaly. (b) Density structure. Numerical value of figure indicates the density (g/cm³).

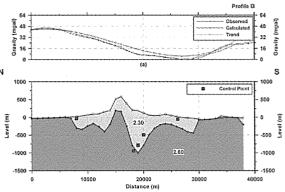


Fig.2 Profile B (a) Gravity anomaly. (b) Density structure. Numerical value of figure indicates the density (g/cm³).

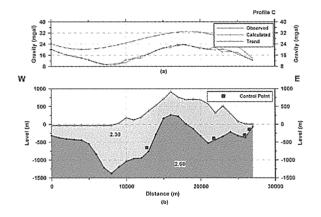


Fig.3 Profile C (a) Gravity anomaly. (b) Density structure. Numerical value of figure indicates the density (g/cm³).

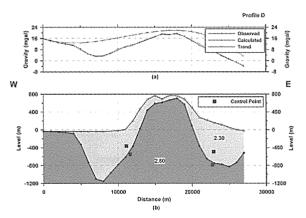


Fig.4 Profile D (a) Gravity anomaly. (b) Density structure. Numerical value of figure indicates the density (g/cm³).

Direct observation of Mg/Si ratios of aqueous fluids in the Earth's upper mantle

T. Kawamoto, K. Matsukage (Ibaraki Univ.), K. Mibe (Geophys. Lab.), M. Isshiki (SPring-8),

K. Nishimura, N. Ishimatsu (Hiroshima Univ.), S. Ono (JAMSTEC)

Direct observation of aqueous fluids coexisting with MgSiO₃ enstatite and/or Mg₂SiO₄ forsterite is performed at 0.5 – 5.8 GPa and 800 -1000 °C with an externally heated diamond anvil cell and synchrotron X-rays. At 1000 °C, forsterite crystallizes below 3 GPa in the MgSiO₃ – H₂O system and not above the pressure. Forsterite congruently dissolves into the aqueous fluids up to 5 GPa at 1000 °C. The present experiments suggest that the aqueous fluids coexisting with enstatite and forsterite have Mg/Si < 1 below 3 GPa and 1 < Mg/Si < 2 above the pressure up to 5 GPa (Figuire). This transitional pressure shows a positive relation with temperature and is located at 2.5 GPa and 800 °C. This chemical variation as a function of pressure should have an effect on the dihedral angles of aqueous fluids that in turn control their movement in the upper mantle. The aqueous fluids that coexist with mantle minerals do have similar Mg/Si ratios to those found in the partial melts of hydrous peridotite. Somewhere within the upper mantle these two fluids unite to form a single regime and cannot be distinguished from each other.

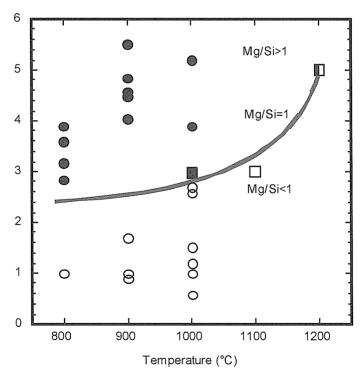


Figure caption Experimental results showing Mg/Si ratios of silicate components in aqueous fluids coexisting with enstatite (En) and forsterite (Fo). The value of 1200 °C at 5GPa and 1000 °C and 1100 °C at 3GPa are after previous studies. At pressures lower than 3 GPa, the aqueous fluids coexisting with En and Fo have Mg/Si <1, and at pressures greater than 3 GPa, they have 1 < Mg/Si <2.

A warning model based on temporal changes of coda Q for volcanic activity at Nevado Del Ruiz Volcano, Colombia

John M. Londono and Yasuaki Sudo

The coda Q has been calculated for Nevado del Ruiz Volcano, Colombia (NRV) from 1985 to 1999 by using a single scattering model. During this period, the inverse of Q (Q-1 proportional to attenuation) exhibited a long-term decrease with time, as well as shorter-term variations related to the volcanic activity. Q-1 increased prior to volcanic crises and decreased afterward. Based on these observations, a seismic warning criterion has been developed. The parameters (frequency band, size of moving average window, and threshold levels) necessary to evidence clear and significant short-term changes in Q-1 have been investigated and appropriated values are proposed. We suggest a phenomenological model with three stages for the short-term temporal changes in Q-1 at NRV. Firstly, Q-1 increases before a volcanic crises because of accumulation of gas and/or liquid, which decreases the aspect ratio of fluid pockets and increases the fractional volume of fluid in the rocks and the pore aspect ratio. Secondly, Q-1 starts to decrease during the crises by the discharging of fluids such as gas, water, etc. from the volcano. Finally, Q-1 becomes more stable after the crisis at a lower value because of the degassing and/or increasing of rigidity of the medium because of the long-term crystallization and cooling processes. Q-1 seems to be a promising monitoring tool at NRV. It is possible that the observed temporal changes of Q-1, combined with other parameters, may help to predict with greater accuracy a volcanic crisis at NRV.

Velocity structure and a seismic model for Nevado del Ruiz Volcano (Colombia)

John M. Londono and Yasuaki Sudo

A seismic tomographic study was performed for Nevado del Ruiz Volcano (NRV) using more than 1500 high-quality local and regional. Three low P wave velocity (low-VP) and low S wave velocity (low-VS) zones were found; one low-VS zone at depths 2?4 km located beneath the volcano; a second low-VP and low-VS zone at depths 5?10 km located beneath the crater, elongated and dipping to the E?SE; and a third low-VP and low-VS zone at 10 to ~12 km farther to the east. These three low-VP and low-VS zones are believed to be the location of heat sources. A high-velocity zone for both P and S waves was found at shallow depths (0~5 km) around the active crater. The upper part (0~2 km depth) of the high-VP and high-VS zone was characterized by low-VP/VS ratios (<1.68), while the deeper part (2~10 km) had high-VP/VS ratios (>1.80). The low-VP/VS zone is correlated with a steam-dominated geothermal system. The high-VP/VS is interpreted as an intrusive body of magmatic origin which includes partial melting zones associated with low-VS anomalies. A small low-VP zone in which long-period (LP) earthquakes were clustered was found to the southwest of the volcano. Based on the data obtained with the tomography in combination with seismicity,

geochemistry, geology and gravimetry, we suggest a model for the seismic activity of NRV. Volcano-tectonic (VT) earthquakes that occur very often in swarm-like patterns located in several clusters around the volcano seem to be due to changes in stress produced by the passing of fluids and/or gas through many small cracks. A fault and a caldera-like structure separate the VT swarms located to the west of the volcano from the source of LP earthquakes.

Source of the Yelagiri and Sevattur alkaline plutons, South India -Examination based on lead isotope ratios-

T. Miyazaki, T. Shibata, M. Yoshikawa and V. Ram Mohan

In the southern part of the Indian Peninsula, there are a number of alkaline plutons of Proterozoic age. In the northern part of the South Indian granulite terrain, the Yelagiri (syenite, pyroxenite) and Sevattur (syenite, pyroxenite, carbonatite) plutons intrude into Archaean epidote-hornblende gneisses. Miyazaki et al. (2000) reported whole rock Rb-Sr isochron ages of 757±32 Ma for the Yelagiri pluton and of 756±11Ma for the Sevattur pluton.

Miyazaki et al. (2003) revealed that the syenitic magma is formed from highly differentiated mantle-derived alkali basaltic magma; the source of this magma is isotopically different from carbonatite. Moreover, they indicated that both the alkaline basaltic and carbonatitic magmas were derived from a subduction-related enriched mantle. This mantle probably occurred at a convergent margin along the southern and/or southeastern edge of the Dharwar craton, and survived, as a sub-continental lithospheric mantle, from the early Proterozoic until at least 750 to 800 Ma ago.

In order to reveal the geochemical characteristics and evolution of this sub-continental lithospheric mantle in detail, we measured the lead isotope ratios of the syenites and examined the source characteristics of the Yelagiri and Sevattur plutons. The ratios ²⁰⁸Pb/²⁰⁴Pb, ²⁰⁷Pb/²⁰⁴Pb and ²⁰⁶Pb/²⁰⁴Pb were in the ranges of 36.919 ~37.751, 15.376~15.437 and 16.562~16.901 in the Yelagiri syenites and 37.115~37.899, 15.374~15.421 and 16.535~16.805 in the Sevattur syenites. In contrast, most carbonatites of the northern part of Tamil Nadu were found to have wider ²⁰⁸Pb/²⁰⁴Pb, ²⁰⁷Pb/²⁰⁴Pb and ²⁰⁶Pb/²⁰⁴Pb ranges: 36.61~38.55, 15.36~15.68 and 16.63~18.62, respectively (Schleicher et al., 1998). Schleicher et al. (1998) indicated two possible explanations of the large lead isotope ratios of the Tamil Nadu carbonatites. One possibility was the mixing of two mantle reservers, and the other was contamination of crustal materials, which have high lead isotope ratios, into magma derived from one mantle reserver. On the other hand, the lead isotope ratios of the Yelagiri and Sevattur syenites show narrow ranges and correspond to the carbonatites with lower lead isotope compositions.

Miyazaki et al. (2003) revealed that the initial neodymium isotope compositions of the Yelagiri and Sevattur syenites [ε Nd = -11.4 ~ -8.1 (760Ma)] were lower than the compositions of the carbonatites [ε Nd = -8.6 ~ -3.8 (760Ma)] in the northern part of Tamil Nadu. They proposed isotopic heterogeneity in the mantle regions, because the lower neodymium isotope composition of the Yelagiri and Sevattur syenites was not explained by contamination of metamorphosed crustal rocks. Therefore, it is considered that the different ranges of the lead

isotope ratios between carbonatite and syenite also reflect isotopic heterogeneity in the mantle.

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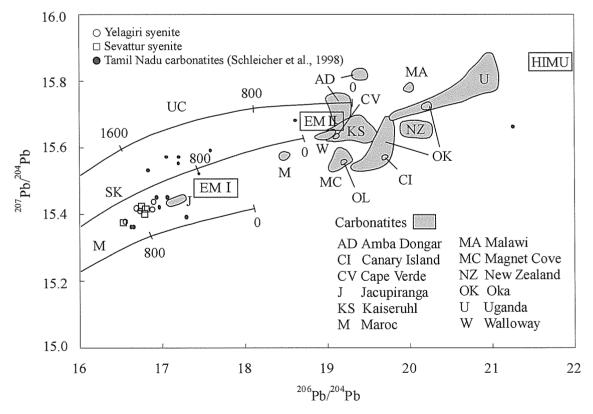


Fig. 1. ²⁰⁶Pb/²⁰⁴Pb vs. ²⁰⁷Pb/²⁰⁴Pb diagram of the syenites and carbonatites from Tamil Nadu.

Influences of the ion-source condition and focusing procedure on reproducibility in a lead isotope measurement by a thermal ionization mass spectrometer (TIMS)

T. Miyazaki, T. Shibata, M. Yoshikawa and B. Vaglarov

Introduction

The lead isotope ratio is an important geochemical tracer. Improving the reproducibility of an isotope ratio measurement is very important to improving the reliability of the isotope analysis. For this purpose, it is critical to carefully establish the optimal measurement methods (e.g., filament material, procedure for sample loading, filament temperature and heating process, condition and tuning procedure for mass spectrometry). Miyazaki et al. (2003a) and (2003b) described about the silica-gel activator and the quality of the Re-filament, respectively.

Setting the conditions and tuning the measurement parameters for mass spectrometry play an important role in influencing the reproducibility of the lead isotope measurement. Special attention should be paid to both the condition of the ion source, which consists of several electrode-lenses and slits, and the focusing procedure for the ion beam through the electrode-lenses. Adhesion of dirt to the ion lenses and imperfect focusing are well known to cause low reproducibility.

To achieve high reproducibility in lead isotope measurements, we examined the effectiveness of ion-source cleaning and strict determination of the focusing procedure.

Experiment

Lead isotope measurements were carried out on a multi-collector mass spectrometer (Triton TI[®] with 9 Faraday cup collectors) at the Beppu Geothermal Research Laboratory, Kyoto University. The solution of NIST SRM 981 (as 100 ng of Pb) was used for this experiment. The preparation method for the lead isotope measurement was that of Miyazaki et al. (2003a). The sample was loaded onto a Re filament (Finnigan[©]; 0.75 mm wide, 0.025 mm thick).

All measurements were carried out by a sequence program, in which all procedures of the measurement, including filament heating and focusing of the ion beam were performed automatically. A filament temperature of approximately 1320°C, which was measured by the pyrometer fitted on the Triton TI®, was used for the measurement (This temperature is equivalent to 1200°C measured by the pyrometer fitted on a MAT262® in our laboratory). The acquisition of data started at approximately 15 minutes after preheating and beam adjustment. The data were obtained in the so-called "static mode," in which the ion beam of each mass number was simultaneously collected by a fixed Faraday cup. These data were computed from 99 repeated measurements, comprising eleven scans in each of nine blocks. The data acquisition took approximately 30 minutes for each measurement.

Cleaning of the ion-source was carried out by wiping with 0.5 N HNO₃, Milli-Q water, and ethanol without disassembling the ion source. After cleaning, the ion source was baked for more than one day by halogen lamps attached in source housing.

Ion source and focusing of the ion beam

The ion source of the Triton TI[®] consists of several electrode-lenses and slits, which are mounted on the ion-source flange. Ions from the filament are accelerated through the high voltage potential and are collimated into a beam by slits in the source electrode-lenses.

Beam-focusing on the Triton TI[®] is fully controlled by a desktop computer. Five parameters of ion lenses (Extraction Left, Extraction Right, Condenser, X-Symmetry, and Z-Focus) and one parameter of wheel position (Wheel Fine Tuning) are controllable. The settings of these parameters must be optimized for each sample and measurement to avoid an additional instrumental bias in the isotope ratios and to ensure maximum reproducibility of measurements. The intensity of a stable ion beam is used as an indicator and is maximized. Each parameter can be manually and automatically adjusted. The parameters of the source lenses can be stored in a hard disk as a computer file. In a sequence measurement, an automatic focusing procedure is ordinarily used, in which several or all parameters are automatically adjusted. The selection of parameters and the order of execution of these parameters are important for perfect focusing and can be determined in the 'Autotune Configuration File'.

Result and discussion

The lead isotope ratios measured with different ion-source conditions and focusing procedures are shown in Fig. 1.

Before cleaning and baking, the dirt of the ion lenses was mainly observed as dark-brownish coating on the surfaces of the electrode-lenses. In particular, slits and some parts of the electrode-lenses, which were composed of the inner portion of the ion lenses, were conspicuously coated by dark-brownish dirt. Weak nitric acid (<0.1 N) was effective in removing this dirt, although the material of this dirt was not identified.

The cleaning and subsequent baking of the ion source were effective in increasing the reproducibility of the isotope ratios. The error of the measurements after cleaning and baking of the ion source was approximately 61% in comparison with the measurements before the cleaning and baking (Fig. 1). The difference in the focusing procedure, which depends on the selection and order of the focusing parameters in the 'Autotune Configuration File', also had a remarkable effect on the reproducibility of the isotope ratios. Adjustment of all parameters and the choice of 'Extraction Symmetry' rather than 'Extraction Left' and 'Extraction Right' in the 'Autotune Configuration File' were necessary to achieve perfect focusing during the sequence measurement. In the present study we determined the perfect focusing procedure with the following parameter selections and order of execution; 1. Wheel Fine Tuning; 2. Extraction Symmetry; 3. Condenser; 4. Extraction Symmetry; 5. Condenser; 6. X-Symmetry; 7. Z-Focus. The error of the measurements with our perfect focusing procedure was approximately 17% in comparison with the measurements obtained with imperfect focusing, in which 'Extraction Left' and 'Extraction Right' were selected instead of 'Extraction Symmetry', and several parameter such as 'Condenser', 'X-Symmetry', and 'Z-Focus' were not adjusted. Moreover, the reproducibility of these results was higher than that of measurements by the MAT262[®] mass spectrometer in our laboratory (Fig. 1).

From the above results, it is concluded that the condition of the ion source and the focusing procedure have a strong influence on the reproducibility of lead isotope ratios. Dirt on the ion source and imperfect

focusing are the cause of low reproducibility with lead isotope ratios. (From this result, it is understood that measurement with a defocused beam is not recommended.) The treatment of the ion source and strict focusing of the ion beam are indispensable to achieving high reproducibility with lead isotope ratios.

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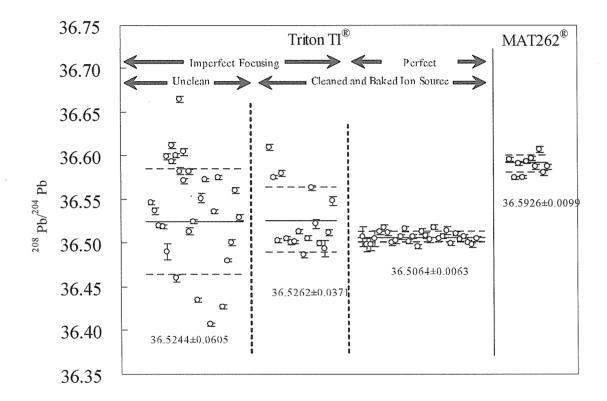


Fig. 1. Variation in $^{208}\text{Pb}/^{204}\text{Pb}$ ratios for sample sizes of 100 ng. The vertical bar with open circles shows the error in $2\sigma_m$ of each measurement. The horizontal solid and dotted lines show the average and range of error (1σ) in the $^{208}\text{Pb}/^{204}\text{Pb}$ ratios.

Identification of multiple faulting of the Yoro fault based on sedimentological core analysis

T. Nakanishi, K. Takemura, M. Togo et al.

A set of close-interval sedimentological analysis and radiocarbon dating of sediment from adjacent core samples across a fault is an instructive tool for identifying multiple faulting events and these timing. Under the Nobi Plain, central Japan, vertical displacement along the Yoro fault is recorded in ca 40 m Holocene sediment. We obtained two core samples from the present delta plain crossing this fault. Sedimentological analyses of lithology, magnetic susceptibility (2 cm-interval), analyses of grain size (10 cm-interval), composition of very fine sand fraction (20 cm-interval), and AMS ¹⁴C measurements (38 samples) were carried out on these core samples. Based on vertical displacement of Kikai-Akahoya ash layer dated at 6,300 yBP, we estimated the average vertical slip rate of the Yoro fault at 1.6 mm/year.

Determination of pre-eruptive water content of silicic magmas using melt inclusions in quartz phenocrysts: examples from the Unzen 1991–1995 dacite

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Museum

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Melt inclusion analysis by FTIR (Fourier transform infrared spectroscopy) can provide information about variation in water content present in pre-eruptive magma. The combination bands at 4500 cm⁻¹ and at 5200 cm⁻¹ are widely used for total water determination. The absorption peak heights of these bands obey the Lanbert-Beer's law and hence are useful to determine water concentrations in melt inclusions. However, molar absorption the coefficients these bands display strong composition-dependence (e.g., Yamashita et al., 1997). We determined the molar absorption coefficients of water in glass of rhyolitic composition and applied the results to natural melt inclusions (rhyolitic glasses) trapped within quartz phenocrysts in Unzen dacite.

Yamashita (1999) synthesized a series of hydrous glasses of rhyolitic compositions (77.4 wt.% SiO_2 under anhydrous condition) and determined the molar absorption

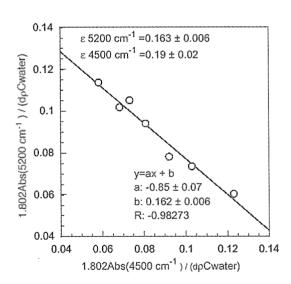


Figure 1 Normalized absorbances of the 4500 cm⁻¹ and 5200 cm⁻¹ bands for synthesized rhyolitic glasses (Yamashita, 1999) using linear baselines. Abs = absorption peak height, d = sample thickness, $\rho =$ sample density and $C_{water} =$ water content. The molar

coefficients at 4500 cm⁻¹ and at 5200 cm⁻¹. We calibrated the molar absorption coefficients for these synthesized glasses by FTIR spectrometer in the Institute for Geothermal Sciences of Kyoto University at Beppu (JASCO 610 Fourier-transform spectrometer which is composed of JASCO Micro 20 IR microscope, InSb detector, CaF₂ beam splitter and W-halogen lamp). The molar absorption coefficients ε of 4500 cm⁻¹ and 5200 cm⁻¹ are determined as 0.19 ± 0.02 and 0.163 ± 0.006 (m³/mol m), respectively (Fig. 1). The present results agree with previous results of Yamashita (1999) ($\varepsilon_{\Box\Box cm-1} = 0.199 \pm 0.009$, $\varepsilon_{5200cm-1} = 0.158 \pm 0.005$) within the analytical errors. With IR spectroscopy using these coefficients, water concentrations up to ~4wt.% in rhyolitic glasses can be determined accurately.

The major element and water concentrations were analysed in melt inclusions from quartz phenocrysts in Unzen dacite. These inclusions have similar SiO₂ contents (> 75 wt.%) to reported values of melt inclusions trapped within plagioclase phenocrysts in Unzen dacite by Holtz et al. (2002). This suggests that the effect of growth of host quartz crystal at the inclusion wall is negligible. These inclusions have a small bubble and are sometimes partially crystallized. In order to determine water concentrations, these inclusions are homogenized by melting and quenching on a heating stage under microscope. Thereafter, the inclusions are doubly polished and analyzed by FTIR spectrometer. Water contents of melt inclusions range between 5 and 7 wt.% H₂O which are slightly out of our standard glass compositions. These values are consistent with reported value of the melt inclusions in plagioclase phenocrysts by Holtz et al. (2002) (6–7 wt.%) by the use of Raman spectroscopy and with results from isotopic studies of Kusakabe et al. (1999) (~6 wt.% H₂O). The H₂O solubility in rhyolitic magmas suggests that the melt inclusions with 5wt.% inclusions were formed at more than 150 MPa and the inclusions with 7 wt.% H₂O were formed at more than 250 MPa.

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GPS observation of crustal movements in the Macolod Corridor, Philippine

T. Ohkura, G.M. Besana, J.A. Daligdig, M. Ando, M. Abigania, E. Mangao, R. Sumanda, R. Esplida and T. Nakano

The origin of the Macolod Corridor (MC) is an unsolved problem, and various hypothesis have been suggested to explain its formation. In order to understand the formation process of the MC, it is very important to determine and understand its present crustal deformation. Therefore, monitoring of the crustal movements in the MC using GPS measurement was initiated as one of an integrated research which includes a paleomagnetic study and K-Ar dating of volcanic rocks by Kyoto University and the Philippine Institute of Volcanology and Seismology (PHIVOLCS).

Thirteen stations were established in 1996 and two stations were added in 1999 in and around the MC. From April 1996 through September 2001, ten GPS campaigns were made. In each campaign, 5 to 8 dual-frequency geodetic GPS receivers (Ashtech Z-XII and Topcon GP-R1DY) were used to acquire the data simultaneously at different stations. Each observation was done continuously for 3 - 5 days and the data were collected at 30 seconds interval for 24 hours.

Using Bernese GPS software Ver. 4.2, a left lateral motion at 2-3 cm/yr with a large amount of counterclockwise motion was detected in the southern part of the study area. Furthermore, a large amount of extension, along NNW-SSE or N-S directions were observed in the MC. These observations are consistent with the block motion after 2Ma determined through a recent paleomagnetic research and a K-Ar dating study of volcanic rocks in and around the MC. It is possible that these crustal movements detected by GPS observations are closely related to the formation process of the MC.

Change in color of thermal water of the Umi-Jigoku hot pool in the Beppu geothermal area

Ohsawa, S., Onda Y. and Takamatsu, N.

In the Umi-Jigoku hot pool in Kannawa area of Beppu City, the acid Cl-SO₄ type thermal water had been colored in *greenish blue*, however we recently observed that the color changed to *blue green*, which is caused by absorption of sunlight of long wavelengths by H₂O molecule (Fig. 1). The previous color (*greenish blue*) results from blending the *blue green* with *blue* caused by Rayleigh scattering of sunlight by aqueous colloidal almino-silicate of 0.025 – 0.45-μm particle size. Owing to low concentration of the aqueous colloid (see Table 1), the color of the Umi-Jigoku thermal water dose not become clear blue in contrast to other blue-colored thermal waters in Beppu. For more information, it is ascertained by calculation of absorbance using dissolved ion concentration and molar extinction coefficient that the thermal water of Umi-Jigoku hot pool is not colored by aqueous ferrous ion. (Submitted to Rep. Oita Prefect. Hot Spring Res. Soc.)

Table 1 Concentrations of aqueous colloids of $0.025-0.45^{\circ} \mu$ m particle size in various blue-colored thermal waters.

Place	Material	0.025~0.45μm (mg dm ^{·3})			
Hotel Kannawa·en	Silica	138			
Kinko-yu *	C.J.	104			
Keikan-yu *	Silica	8.0			
Kamado-Jigoku	Silica	27.0			
Umi-Jigoku (Mar. 9, 2002)	A1 ' '1' '	0.57			
Umi-Jigoku (Jan. 7, 2003)	Almino-silicate	0.08			

^{*}Restaurant Ichinoide-kaikan

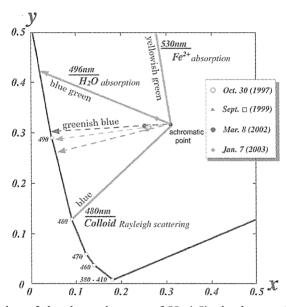


Fig. 1 Change in color of the thermal water of Umi-Jigoku hot pool shown on chromaticity diagram of Yxy color space.

Seismic Activity around Aso Caldera and Kuju Volcanoduring the period from April in 2002 to March in 2003

Y. Sudo, T. Ohkura and S. Yoshikawa

A telemetrically seismic network around Aso Caldera and Kuju Volcano is composed of 17 stations. Several earthquakes have occurred at the line from the north side area of Kuju Volcano to the west side area of Aso Caldera.

The epicentral cluster near Mt.Nakadake crater was practically the origin of isolated volcanic tremor. The other cluster near Tateno Valley just outside of the west caldera rim was the swarm occurred on 13 February 2003. The largest event occurred at 11:27 on 13 February with magnitude 2.2. The mechanism of this

event may indicate a combination of both normal and right lateral strike-slip type faults. The tension axis maintains a horizontal north-south orientation. In Aso Caldera, there was no remarkable seismic event during this period.

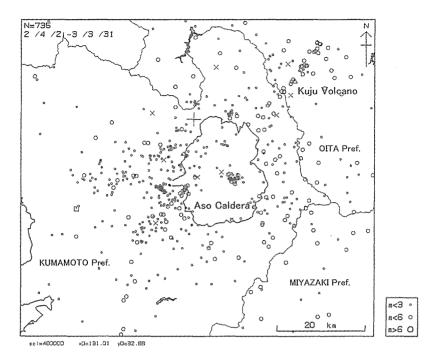


Fig. 1. Epicenteral distribution of events occurred at the region of Aso Caldera and Kuju Volcano during the period from April in 2002 to March in 2003. A trianlgular mark indicates Mt.Kuju in Kuju Volcano and Mt.Nakadake in Aso Volcano. A cross indicates aseismic station. A square ndicates the Kumamoto city office.

Lake Biwa sediments in Japan: Recorders of global paleoenvironment and island arc tectonics since Pliocene

K. Takemura, A. Hayashida, T. Nakagawa et al.

Lake Biwa is the largest freshwater lake in Japan, measuring 22.6 km side by 68 km long and having a maximum depth of 104 m. The lake has a long history from the early Pliocene based on the geological survey and paleogeographical study of lake sediments around Lake Biwa (Kobiwako Group). The sediments revealed the paleoenvironmental change and tectonic events at the convergent margin of the Eurasian plate since early Pliocene times.

Deep drillings for paleolimnological studies were carried out 1970's and 1980's. These studies showed that the present Lake Biwa Basin bears a sedimentary sequence of about 900 m thickness, which were deposited in lacustrine or fluvial environments in the Pliocene and the Quaternary. The upper most unit is a continuous and massive sequence of lacustrine clay of 250m thick intercalating more than 50 layers of volcanic ashes. The tephrochronological and paleoclimatological data suggest that the 250 m clay unit has been deposited from the last 430 kyrs and is correlated to major glacial-interglacial cycles.

Pioneer studies on the long sediment core from Lake Biwa revealed that the fluctuations of climatic

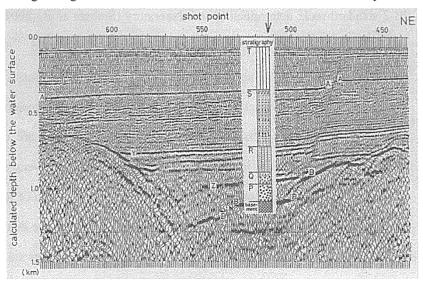
signals were successively correlated to the Milancovitch cycles. However, the upper part of long core was highly disturbed because of relatively poor coring techniques. In 1990's, 15 to 20 m long piston cores were obtained from three sites in Lake basin, and analytical studies at higher-resolution have been carried out. These studies suggested that the Lake Biwa sediments record climatic fluctuations at millennial to century order, which can be correlated to the Heinlich events or Dansgaard-Oeschger cycles. The paleolimnological setting of Lake Biwa offers outstanding advantages for high-resolution study of the East Asian monsoon climate of the Late Quaternary.

Active tectonics in central Japan and the significance of sedimentary record from Lake Biwa

K. Takemura, N. Inoue and T. Ishiyama et al.

Deep seismic reflection survey and gravity measurement in Lake Biwa through 1970's and early 1980's brought the information on the basement topography, and revealed that tilting structure from east to west. However, the deepest part is located in the Central Basin where Prof. Horie's team drilled in early 1980's from the seismic reflection survey, and gravity survey indicated that the lowest Bouger anomaly area is north Lake Basin. Seismic reflection survey revealed the existence of active fault in the deep basin. The lithostratigraphical information indicated the subsidence history of the present lake basin during about 2.5 my. Shallow seismic reflection surveys carried out in 1980's and 1990's show the distribution of active fault traces and size of displacement along the coast of Lake Biwa. Taking an additional core from the Northern part of the basin will allow us to look out the difference between two parts of the basin. Chronological control will be carried out from the tephrochronology, magnetic measurement and establishing pollen stratigraphy. If the time recovery is complete, the result will bring out good continuous information of the climatic history and

subsidence history of the basin. Even incomplete time coverage will still bring the information about the tectonic disturbance of the basin. Turbidite layers, which will be detected from Lake Biwa sediments, will also bring the information about recent earthquake event (Historical earthquake) and periodicity.



Multichannel seismic reflection profile along Line 9-1 in Lake Biwa. Core site was indicated by arrow. Inset roughly correlates reflectors with stratigraphic column.

Sr and Nd isotopic ratios and Rb, Sr, Sm and Nd concentrations of JB-2 rock reference material with isotope dilution method

Masako Yoshikawa and Tomoyuki Shibata

The Rb-Sr and Sm-Nd systematics are basic tool to obtain age information of geological samples. We describe the analytical procedure and report the results for Sr and Nd isotopic ratios and Rb, Sr, Sm and Nd concentrations of JB-2 (basalt) Geological Survey of Japan (GSJ) standard rock to assess the reliability of natural samples employed with isotope dilution (I.D.) method.

The analytical procedure for chemical separation and mass spectrometry basically followed Yoshikawa et a.l (2001) and Shibata et al. (2001). Rb, Sr, Sm and Nd concentrations were determined by I.D. method using ⁸⁷Rb, ⁸⁴Sr, ¹⁴⁹Sm, and ¹⁵⁰Nd spikes, respectively. ⁸⁷Sr/⁸⁶Sr and ¹⁴³Nd/¹⁴⁴Nd ratios were estimated with determination Sr and Nd concentrations, respectively. Mass spectrometry was carried out on a Finnigan MAT 261. To correct for isotopic fractionation for unspiked ⁸⁷Sr/⁸⁶Sr and ¹⁴³Nd/¹⁴⁴Nd measurement, normalizing factor for Sr and Nd are ⁸⁶Sr/⁸⁸Sr = 0.1194 and ¹⁴⁶Nd/¹⁴⁴Nd = 0.7219, respectively. Measured ratios for standard materials were ⁸⁷Sr/⁸⁶Sr = 0.710279 \pm 0.000028 (2 σ) for NIST SRM987 (n = 5) and ¹⁴³Nd/¹⁴⁴Nd = 0.511851 \pm 0.000013 (2 σ) for La Jolla (n = 9).

Analytical results are listed in Table 1. The mean values for spiked 87 Sr/ 86 Sr and 143 Nd/ 144 Nd ratios with reproducibility (2sm) are 0.703734 ± 0.000007 and 0.513084 ± 0.000015 , respectively. This is consistent with unspiked 87 Sr/ 86 Sr (0.703703 ± 0.000011) and 143 Nd/ 144 Nd (0.513097 ± 0.000011) ratios of JB-2 standard rock, determined by our laboratory (Shibata et al., 2001). Our results are in good agreement with the averages (87 Sr/ 86 Sr 0.703700 ± 0.0000105 , error is 2σ) and 143 Nd/ 144 Nd (0.513080 ± 0.000024 , error is 2σ) of the compiled data (Shibata et al., 2001) except Nohda and Wasserburg (1981).

The mean concentrations of Rb, Sr, Sm and Nd concentrations are 6.65, 177, 2.27, and 6.31, respectively. The data obtained by the I.D. method for JB-2 are compiled in Table 2 for comparision. Although Rb concentration in this study is 9% higher than the averaged values of compiled data, our results are consistent with the compiled data within error. The variation of Rb concentrations was larger than those of other elements between the compiled data (Table 2). It may be due to heterogeneous concentrations of Rb in this rock. From these observations, it is indicated that our results are accurate and this methods is reliable.

Table 1. Analytical results of Rb, Sr, Sm and Nd concentrations (μg/g) and Sr and Nd isotopic ratios of JB-2.

	Rb	Sr	Rb/Sr	⁸⁷ Sr/ ⁸⁶ Sr	$2\sigma_{\rm m}$	Sm	Nd	Sm/Nd	¹⁴³ Nd/ ¹⁴⁴ Nd	$2\sigma_{\rm m}$
	6.50	177.1	0.03669	0.703720	0.000008	2.27	6.253	0.3630	0.513088	0.000017
	6.63	176,4	0.03757	0.703703	0.000011	2.26	6.352	0.3563	0.513068	0.000019
	6.73	176.7	0.03806	0.703699	0.000011	2.27	6.301	0.3609	0.513102	0.000018
	6.65	176.9	0.03761	0.703712	0.000017	2.26	6.283	0.3597	0.513077	0.000017
mean	6,65	176.8	0.03763	0.703709	0.000007	2.267	6.312	0.3600	0.513084	0.000015
RSD (%)	1.4	0.18	1.52			0.28	0.66	0.78		

Table 2. Compiled Rb, Sr, Sm and Nd concentrations employing the isotope dilution method and Sr and Nd isotopic ratios for JB-2 rock standard.

	Rb	Sr	Rb/Sr	⁸⁷ Sr/ ⁸⁶ Sr	error	Sm	Nd	Sm/Nd	143 Nd/ 144 Nd	error	Refernce
	6.51	182	0.03577	0.703450	0.00003	2.29	7.46	0.3070	0.513110	0.000020	1
	5.17			0.703753	0.000014						2
	6.2	173	0.03584	0.70375							3
						2.27	6.37	0.3564			4
						2.24	6.22	0.3601	0.513090	0.000030	5
						2.242	6.312	0.3552			6
						2.24	6.22	0.3601			7
						2.26	6,305	0.3584	0.513089	0.000030	8
	6.524	180.3	0.03618			2.31	6.4	0.3609			9
	6.33										10
	6.3	176	0.03580	0.70376						*******************************	11
average	6.105	176.4	0.03594	0.703754	0.000010	2.260	6.305	0.3585	0.513090	0.000024	
RSD (%)	8.8	2.1	0,6			1.2	1.2	0,6			

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^{3.} Kurasawa (1984)

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^{5.} Tanaka (HP of GSJ) 6. Nakamura (HP of GSJ)

⁷ Kamioka and Tanaka (1990)

^{8.} Pin and Zalduegui (1997)

^{9.} Yuhara et al., (2000)

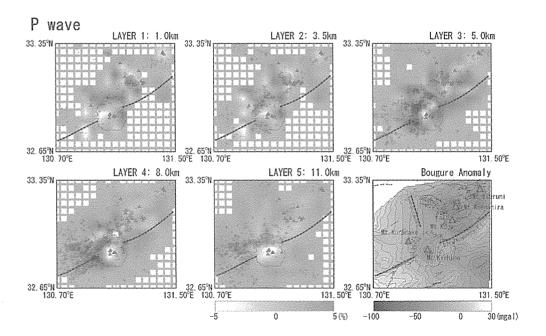
^{10.} Togashi (HP of GSJ)

^{11.} Zichao (HP of GSJ)

Three-dimensional seismic velocity structure of the Hohi Volcanic Zone in central Kyushu, Japan

M. Yoshikawa, Y. Sudo, J. M. Londono and H. Masuda

We studied the three-dimensional seismic velocity structure of the Hohi Volcanic Zone (HVZ) in central Kyushu, Japan, using the tomographic method of Zhao et al. (1994) for exploring the detailed volcanic and geothermal structure. Linear regions with low Vp at depths of 5 and 8 km are located, which lie along the volcanic chain within the HVZ. At these same depths belts of high Vp and high Vs run parallel to the Oita-Kumamoto tectonic line to the east and west of Aso Caldera, while at a depth of 11 km a low Vs belt runs parallel to the Oita-Kumamoto tectonic line. This low Vs belt is connected to low Vp and low Vs regions near Mt. Kishimadake in Aso Volcano and the south of Kuju Volcano at the same depth of 11 km. These seismic velocity anomalies are related to the seismic location, the temperature distribution and the Curie point depth. The seismic velocity structure in this study can thus be linked to other investigations of the volcanic and geothermal structure of the HVZ.



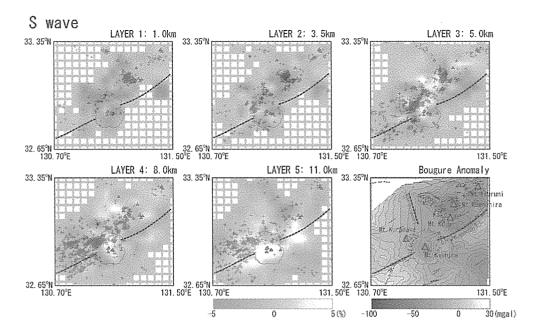


Fig. 1 Plan views of perturbations of P wave and S wave velocities rounding off $\pm 5\%$, hypocenter distributions (+), the distribution of faults (The research group for active tectonics in Kyushu 1989) and the Bouguer anomaly (Gravity Research Group in Southwest Japan 2001). Carrying out this inversion on the study volume which consists of the region 32.6° N -33.5° N x 130.6° E -131.5° E wide and from 0 to 30 km depth, we adopt a 0.05° spacing of nodes in horizontal planes at Z = 1.0 km, 3.5 km, 5.0 km, 8.0 km, 11.0km and 20.0 km depth, Z = 0 km indicates sea level.

Holocene Environmental Changes from Coastal Lakes along the Eastern Coast of Korea

J.G. Yum, K.M. Yu, K. Takemura, et al.

Lithologic, geochemical, and paleontological data from eleven cores (three boring cores, three gravity cores and five vibrocores) of the Hwajinpo and Songjiho lagoons provided the evolutionary histories of each lagoon related to Holocene sea level changes of the East Sea (Japan Sea). The Hwajinpo and Songjiho lagoons located 25 km apart on the mid-eastern coast of Korea, where are known tectonically stable and their similarity and simultaneity present the regional environmental change. Grain size analysis, water content analysis, magnetic susceptibility, and soft X-ray analysis of core samples were used to reconstruct sedimentary environments, as were total organic carbon, total nitrogen, C/N, total sulfur and C/S chemical records.

Assemblages of mollusc and palynomorphs remains also provided paleoenvironmental information.

The morphological classification with the measurement of notches and benches around the Hwajinpo and Songjiho lagoons provide the past erosion processes related to sea and/or lake level change history. There are two major erosion landforms (about 80 cm and 160 cm above mean sea level) in the study area. The each erosion landform could be related with high standing sea level span during the Holocene. The reconstruction of the paleoenvironment from this study allows the establishment of evolutionary models and sea level tends of both lagoons.

公表論文および著書 Publications

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- 高木憲明・金嶋聡・森健彦・山本希・川勝均・大倉敬宏・須藤靖明:地震計アレイを用いた 阿蘇山・火山性連続微動の震源決定、地球惑星科学関連学会2002年合同大会,東京2002.
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- Program) International Workshop "Lake Biwa and Lake Suigetsu: Recorders of Global Paleoenvironments and Island Arc Tectonics" (Kyoto University, November 2002)
- 田中良和・橋本武志・小野博尉・北田直人・吉川慎・松島健、阿蘇火山における空中磁気測量ー繰り返し観測に向けて一、平成14年度京都大学防災研究所研究発表講演会、平成15年2月21日京都リサーチパーク
- 田中良和・橋本武志・小野博尉・北田直人・吉川慎・松島健,阿蘇火山における空中磁気測量 一繰り返し観測に向けて一,京都大学防災研究所研究発表講演会(京都),Feb. 20-21, 2003.
- 田中良和・橋本武志・北田直人・松島 健、阿蘇火山における空中磁気測量—繰り返し観測 に向けて—、KS-What 研究会、12/14、九州大学国際研究プラザ
- 田中良和・橋本武志・北田直人・大久保綾子・宇津木充・清水洋・松島健、雲仙火山及び阿蘇火山における空中磁気測量について、平成14年度CA研究会、12/25、京都大学防災研究所
- 田中良和、地磁気変化を利用した地熱の推移に関する研究(2)、大分県温泉調査研究会、7/31、 別府市鶴見荘
- 鄭国東・*鈴木勝彦*・徐勝・高野 穆一郎・久野 章仁・松尾 基之・史基安・王其,青海湖 堆積物コア試料の地球化学的特徴および古堆積環境の変化,2002 年度日本地球化学会
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- 宇都智史・吉田賢一・塩崎一郎・大志万直人・笠谷貴史・鍵山恒臣・*橋本武志*・中尾節郎・ 矢部征,大山火山の深部比抵抗構造調査(序報),地球惑星科学関連学会(東京),May 27-31,2002.
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- 吉川美由紀・須藤靖明・J.Londono・増田秀晴: 九重火山とその周辺のP波速度構造地球惑星科学関連学会 2002 年合同大会,東京 2002.
- 柚原雅樹・**宮崎** 隆・加々美寛雄(2002)伊那領家帯,吉瀬花崗閃緑岩の年代学的,同位体岩石学的研究.地球惑星科学関連学会 2002 年合同学会(東京, 2002 年 5 月).
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- 吉川美由紀・須藤靖明・J.M.Londono・増田秀晴・吉川 慎: 阿蘇火山および九重火山で見られた地震波速度異常、京大防災研究所研究発表講演会、京都市 2003.
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共同研究 Collaborations

国内

大沢信二 別府海地獄の地球化学的・色彩学的研究 参加機関:東邦大学理学部化学科

大沢信二 別府地熱地帯に産する酸性泉中の希土類元素の存在度 参加機関:東邦大学理学 部化学科

川本竜彦:東京大学地震研究所 一般共同研究、研究員

川本竜彦:愛媛大学地球深部ダイナミクス研究センター 客員研究員

川本竜彦:岡山大学固体地球研究センター 嘱託研究員

川本竜彦:スプリング8高輝度光科学研究センター 外来研究員

柴田知之:同位体分析に基づく島弧マグマの起源並びに進化過程の解明(JAMSTEC, 巽、芳川)

柴田知之:ハワイ・ハレアカラ火山沖の海底火山帯の地球科学的研究(東工大・高橋・芳川)

須藤靖明: 阿蘇火山火山性微動観測ASO2002 (震研・東工大,川勝・山本・金嶋・森・高木・須藤・大倉・吉川慎・迫、2002 年 5 月 14 日 - 16 日:9 月 20 日 - 22 日:11 月 5 日 - 11 日: 2003 年 2 月 25 日 - 28 日)

須藤靖明・大倉敬宏・吉川慎 北海道駒ヶ岳における火山体構造探査(北大、東大など多数)

須藤靖明・大倉敬宏・吉川慎 富士山周辺における地震観測網の構築(東大、京大防災研など多数)

竹村恵二:「京都盆地の地盤に関する研究」関西地盤情報活用協議会

竹村恵二:「大阪湾海底地盤に関する研究」大阪湾海底地盤情報活用協議会

竹村恵二:「京都盆地の構造と地盤災害ポテンシャル評価に関する研究」(京都大学防災研究 所:三村 衛)

田中良和・橋本武志:「雲仙火山のTDEM調査」(東大地震研、京大防災研)

橋本武志:雲仙科学掘削

橋本武志:ロングバレーカルデラ電磁気観測(USGS)

国際

竹村恵二:「High resolution reconstruction of Holocene environmental changes from lagoonal sediments along the East Sea (the Japan Sea)」(Prof. Yu, Yonsei University, Korea)

定常観測 Routine Observations

Geophysical Monitoring Under Operation at AVL

Aso Volcanological Laboratory

Permanent Stations

Nakadake monitoring network

Seismic Stations: HNT, PEL, KSM, SUN, KAE, KAE, KAN, UMA, TAK (microwave telemetry)

Tiltmeters: HNT (water tilt 3-comp.), SUN, KAE, NAR, UMA, KAK (on-site logging)

Extensometers: HNT (invar 3-comp.)

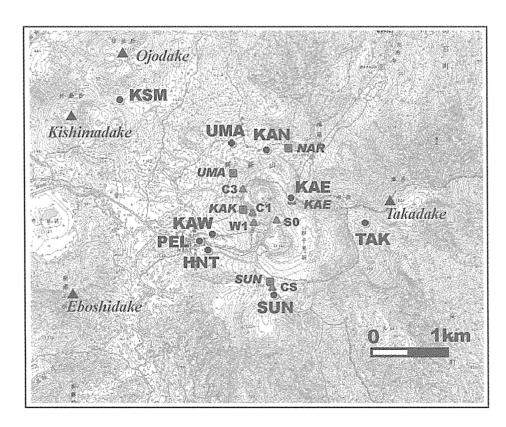
Microphone: HND (microwave telemetry)

Geomagnetic Stations: C1, C3, S0, W1, CS, NGD, FF1 (proton; on-site logging)

C223 (fluxgate 3-comp.; on-site), newC223 (fluxgate 3-comp.; online)

FF2 (proton; online)

Ground Temperature: KAK (boreholes of 70 and 150 m deep; microwave telemetry)



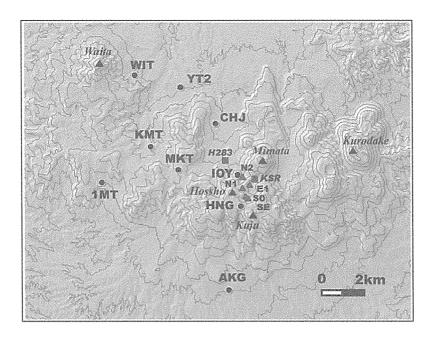
Seismic, geodetic and geomagnetic stations in the central part of Aso.

Kuju monitoring network

Seismic Stations: HNG (radio-telemetry), AKG, CJB, IOY (on-site logging)

Tiltmeters: H283, KSR (on-site logging)

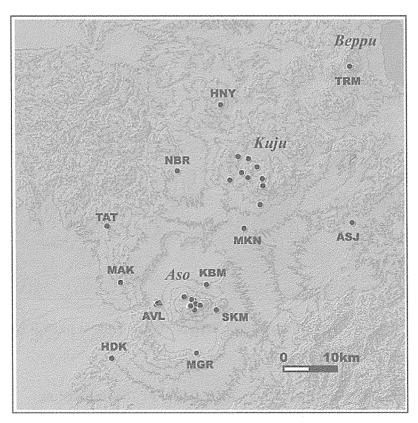
Geomagnetic Stations: N2, E1, S0, SE (proton; on-site logging)



Seismic, geodetic and geomagnetic stations in Kuju area.

Central Kyushu regional network

Seismic Stations : AVL(6), MAK, NBR, MKN, HDK, TAT, MGR (online telemetry) ASJ, HNY, SKM, KBM, TRM (dial-up)



Seismic network in the central Kyushu.

装置 Instruments

【別府】

ICP 発光分光分析装置

波長分散型電子プローブマイクロアナライザー (海洋科学技術センターに貸し出し中)

エネルギー分散型電子プロープマイクロアナライ ザ

波長分散型蛍光 X 線分析装置 エネルギー分散型蛍光 X 線分析装置 粉末 X 線回折装置

液体シンチレーションシステム イオンクロマトグラフ

【阿蘇】

阿蘇, 九重火山連続地震観測システム 地殻変動観測坑道 孔中温度観測システム ビデオ映像監視システム プロトン磁力計 フラックスゲート磁力計 地磁気絶対測定システム 傾斜計 ガスクロマトグラフ

自動滴定装置

ピストンシリンダー型高圧発生装置

ICP-MS 用レーザーアブレーション装置

四重極型 ICP-MS 装置

表面電離型質量分析装置

外熱式ダイアモンドアンビル

フーリエ変換型近赤外分光光度計

赤外顕微鏡

加熱ステージ

可搬型地震計(広帯域,短周期)

人工震源車

重力計

超伝導重力計

地磁気地電流測定装置(広帯域型 ULF, ELF, VLF

型)

光波測距儀

水準測量システム(自動読み)

[Beppu]

ICP emission Spectrometer

Wavelength dispersive electron microprobe (lent to JAMSTEC)

Energy dispersive electron microprobe analyzer

Wavelength dispersion type X-ray Fluorescence analyzer

Energy dispersionty type X-ray Fluorescence analyzer

Powder X-ray diffractometer

Liquids scintillation system

Ion chromatography

Gas chromatography

Automatic titration system

Piston cylinder type high pressure apparatus

Laser ablation system

Inductively coupled plasma mass spectrometer

(ICP-MS)

Thermal ionization mass spectrometer (TIMS)

Externally heated diamond anvil cell

FT-NIR spectrometer

IR microscope

Heatings stage

[Aso]

Continuous seismic monitoring system for Aso and

Kuju Volcanoes

Observation tunnel for ground deformation

Borehole temperature monitoring system for Aso

Video monitoring system of Aso and Kuju Volcanoes

Proton and fluxgate magnetometers

Geomagnetic absolute measurement system

Tiltmeters

Portable seismometers (broadband short period)

Car-mounted seismicsource

Gravimeters

Super-Conducting Gravimeter

Magneto-Telluric measurement system (broad-band

type, ULF, ELF, VLF-band)

Electronic distance measurement system

Leveling survey system (automatic reading)

設備 Facilities

岩石粉砕·鉱物分離室

パックミル・ディスクミルによる岩石粉砕やアイソダイナミックセパレータによる鉱物分離を行う.

器具洗浄室

実験に用いる器具の洗浄を行う. ドラフト2台・イオン交換筒・Milli-Q が設置されている.

クリーンルーム

HEPA フィルターを設置し極力金属使用を控えた設計で、クラス 100 のクリーン度を達成している. Sr・Nd・Pb 同位体比分析のための化学処理(試料の分解・イオン交換クロマトグラフィーによる目的元素の抽出)を行っている.

地下観測坑道(阿蘇火山地殼変動観測坑道)

阿蘇中岳第一火口から南西 lkm の,地下 30m に設けられた,直角三角形の水平坑道で,1987 年度に竣工 した.現在は,水管傾斜計(25m),伸縮計(20,25m),短周期地震計,長周期地震計,広帯域、地震計、 強震計、超伝導重力計が設置されている.

火山研究センター構内地震観測システム

火山研究センター構内では、従来からトリパタイトによる地震観測を行ってきたが、平成13年度に、ノイズ低減の為、約200mのボーリング孔を4本掘削し、孔底に地震計を導入した。これにより、S/N比は大幅に改善され、従来識別できなかった中岳の長周期微動が検出されるようになった。また、ボーリングコアを採取したことにより、研究センターの丘、高野尾羽根(たかのおばね)火山について地質学的に新たな知見が得られつつある。これは、阿蘇中央火口丘の噴火史を研究する上でも貴重な資料である。

研究費 Funding

科学研究費補助金

- 基盤研究(A) 大倉敬宏(分担) 「地震の準備過程の解明ー南アフリカ金鉱山における地震包囲網の完成ー、代表: 飯尾能久 | 33,400 千円
- 基盤研究(B) 竹村恵二(分担)「大規模堆積平野縁の第一級活断層に伴う変位地形と地下構造に関する研究, 代表:岡田篤正」1,700 千円
- 基盤研究(C)大沢信二(代表)「水文化学的手法による火山湧水からのマグマ起源二酸化炭素の流出フラックスの評価」1700千円
- 若手研究(B) 川本竜彦「地球深部条件におけるマグマの分光学」600千円
- 若手研究(B) 橋本武志「自然電位および土壌ガスの確率的トモグラフィによる火山の浅部熱水系の解明」 2,650 千円
- 特定領域研究(2) 田中良和(代表)・須藤靖明(分担)「火山爆発の発生場と発生過程」5,100千円

奨学寄附金等

川本竜彦、日産科学振興財団「富士火山の下でのマグマと H20 フルイドの間の超臨界現象」1800 千円

受託研究

- 竹村恵二 科学技術振興調整費 雲仙掘削 雲仙地溝の活構造解析 1,566 千円
- 田中良和(代表)・橋本武志(分担)科学技術振興調整費「科学掘削による噴火機構とマグマ活動解明のための国際共同研究」8,000千円
- 田中良和・橋本武志(分担) 科学技術振興調整費「雲仙火山: 科学掘削による噴火機構とマグマ活動解明の ための国際共同研究、代表: 宇都浩三」の分担課題: 火山流体の分布と移動に関する電磁気学 的モデルの構築、4,955 千円

共同研究等

- 大倉敬宏、地震研究所一般共同研究「芸予地震(2001年3月24日はスラブマントル内地震か?」128千円 川本竜彦、東京大学地震研究所「一般共同研究」419千円
- 竹村恵二 京都大学防災研究所一般共同研究 (担当者:三村 衛)870千円
- 竹村恵二 ICDP International workshop "Lake Biwa and Lake Suigetsu: Recorders of Global Paleoenvironments and Island Arc Tectonics" 30,000US\$
- 竹村恵二 経済産業省原子力保安院 「地表断層の形状の三次元化及び高精度化と深部起震断層との相関」 (代表:岡田篤正) 8,369 千円
- 田中良和(代表)・橋本武志(分担)日米科学協力事業「火山活動にともなう電磁場の発生機構-阿蘇とロングバレーの比較研究-」 1,567千円

教育活動 Education

学位・授業 Academics

			学位・授業	Academics
学位審查				
大倉敬宏	(審査員)	本多茂喜	(修士	京都大学理学研究科)
大沢信二	(審査員)	網田和宏	(博士	京都大学理学研究科)
	(審査員)	恩田裕二	(修士	東邦大学理学研究科)
須藤靖明	(審査員)	網田和宏	(博士	京都大学理学研究科)
	(審査員)	Estu Kriswa	ti (修士	京都大学理学研究科)
	(審査員)	Sri Hidayati	(修士	京都大学理学研究科)
田中良和	(審査員)	二穴喜文	(博士	京都大学理学研究科)
	(審査員)	細川敬介	(博士	京都大学理学研究科)
	(審査員)	山下 太	(博士	京都大学理学研究科)
	(審査員)	網田和宏	(博士	京都大学理学研究科)
	(審査員)	中坊 真	(博士	京都大学理学研究科)
	(審査員)	金子雅裕	(修士	京都大学理学研究科)
	(審査員)	松岡洋介	(修士	京都大学理学研究科)
竹村恵二	(審査員)	網田和宏	(博士	京都大学理学研究科)
	(審査員)	杉戸信彦	(修士	京都大学理学研究科)
	(審査員)	谷美由起	(修士	京都大学理学研究科)
由佐悠紀	(主查)	網田和宏	(博士	京都大学理学研究科)

講義・ゼミナール

科目 担当教官

(学部)

地熱学 由佐悠紀・竹村恵二・大沢信二・古川善紹

地殼物理学 竹村恵二

陸水物理学 由佐悠紀・諏訪浩1

火山物理学 田中良和・須藤靖明・古川善紹・大倉敬宏

課題演習 D2 尾池和夫²·中西一郎²·久家慶子²·柳谷俊¹·渋谷拓

郎」・古川善紹・大倉敬宏

課題演習 D3 地下構造と活構造・地表変動 岡田篤正2・竹村恵二・堤浩之2・赤平純平・福岡浩・

岩田知孝

課題研究 T8 地表変動・固体地球物理・火山物理 岡田篤正2・竹村恵二・須藤靖明・堤浩之2・入倉孝次

郎・佐々恭二

考古学特殊講義(文学部) 竹村恵二

(大学院修士課程)

第四紀地質学 竹村恵二

水圏地球物理学 II 由佐悠紀・大沢信二・諏訪浩 1 由佐悠紀・世中良和・大沢信二 由佐悠紀・田中良和・大沢信二

地球熱学・地熱流体学 II 須藤靖明・古川善紹・大倉敬宏・竹村恵二

応用地球電磁気学

応用地球科学ゼミナールⅠ

応用地球科学ゼミナールⅠ

大志万直人・田中良和

岡田篤正・竹村恵二・堤浩之2・小松原琢

岡田篤正・竹村恵二・堤浩之2・小松原琢

地球惑星科学特殊研究 (修士論文)

(大学院修士課程および博士後期課程)

水圏地球物理学ゼミナールIII

地球熱学・地熱流体学ゼミナールⅠ

由佐悠紀・大沢信二・諏訪浩¹・斎藤隆志¹

由佐悠紀・田中良和・大沢信二・川本竜彦・鈴木勝彦・

柴田知之

固体地球物理学ゼミナール IV

地球熱学・地熱流体学ゼミナールII

尾池和夫²·中西一朗²·大倉敬宏·加藤護³

須藤靖明・古川善紹・大倉敬宏・小野博尉・橋本武志・

竹村恵二

応用地球科学ゼミナールⅠ

応用地球科学ゼミナール I 地球牛物圏史セミナー 岡田篤正2・竹村恵二・堤浩之2・小松原琢

岡田篤正²·竹村恵二·堤浩之²·小松原琢

増田富士雄3・前田晴良3・竹村恵二・大野照文4

(1)防災研究所、2地球物理学教室、3地質学鉱物学教室、4総合博物館)

野外実習

地熱学野外実習(8月4日~7日)

D2 阿蘇実習(8月1日~8月4日)

応用地球科学現地実習(6月1日~6月4日)

講義(他大学)ほか

北田直人 大阪女子大学:非常勤講師

竹村恵二 集中講義:岡山理科大学総合理学特別講義 VI

大沢信二 集中講義:大分大学教育福祉科学部環境科学入門

セミナー Seminars

地球熱学ゼミナール

2002年

4月 24日 宮崎隆 「火成炭酸塩岩と堆積成炭酸塩岩のミキシング スリランカEppawala岩体の例」

5月 15日 任鐘元 「Petrological study of the Submarine Haleakala East Rift Zone, Hawaii」

5月 22日 鈴木 勝彦「Osmium isotope composition: Global and vertical perspective」

6月 5日 竹村恵二「別府湾周辺の活断層:分布・活動度・活動時期」

6月 26日 大沢信二「雲仙火山におけるマグマ起源二酸化炭素の浅層地下水系への逃散」

7月 3日 網田和宏「火山体内における火山ガスと土壌ガスの混合プロセス」

9月 25日 芳川雅子 「Rb-Sr and Sm-Nd isotopic systematics of the Hayachine-Miyamori ophiolitic complex:

New constraints to origin and Petrogenesis

10月 16日 山田誠 「「河床下の地下水流動」」

10月 23日 網田和弘「「マグマ性ガスと地下水・空気の混合に関する研究」」

- 11月 6日 Chang·Hwa Chen (陳 中華) and the IMAGES Tephra Research Group 「Climatic and Human Impact of Explosive Volcanism」
- 11月 27日
- 12月 4日 新名真由美 「別府地熱地域 伽藍岳の地熱構造モデル―ニ相流を中心とした解釈・」
- 12月 18日 Chang-Hwa Chen (陳 中華)and S. Nakada 「Insights into magmatic variations of Unzen: evidence from isotopic study of Phase I Drilling」

2003年

- 1月 15日 西村光史 「珪長質マグマ溜まりの固結前線における沈降結晶の捕獲」
- 1月 22日 網田和弘 「火山地域から放出される噴気ガスの形成機構に関する研究」

特別セミナー・BGRL セミナー

2002年

- 4月 30日 羽生「ハフニウム同位体を用いた島弧マグマ起源の研究」
- 6月 19日 Kevin Jhonson 「Plume ridge interaction on Southeast Indian ridge」
- 6月 21日 高橋栄一 「Melting process in the Hawaiian plume 」
- 7月 2日 伊藤雅史 (JAMSTEC) 「The winter monsoon fluctuation in eastern Asia over the past 180 ky」
- 7月 10日 Santos 「Re-Os systematics of Palawarn chromititics complex in Phillipin.」
- 7月 15日 辻村知之(岡山大学固体地球研究センターD2) 「硫黄を含むケイ酸塩ガラスのネットワーク縮重合と溶存硫黄種に関する分光学的研究」
- 9月 10日 辻村知之(岡山大学固体地球研究センターD2) 「ラマン分光法によるケイ酸塩ガラスネットワーク 構造に対する硫黄の効果と溶存硫黄種の関係」
- 12月 9日 真島英壽(九州大学地球惑星科学科 研究生) 「海洋島型玄武岩の起源-九州北西部玄武岩類を例として」
- 12月 17日 松山明人氏(国立水俣病綜合研究センター) 「中国貴酬省有機化学工場排水による水銀汚染について」
- 12月 24日 高橋俊郎(海洋科学技術センター固体地球統合フロンティア研究システム分析解析センター):富山県南部、中新統岩稲累層を構成する多様な安山岩類の成因

AVLセミナー

2002年

- 4月 16日 古川邦之「流紋岩質マグマにおけるスパター噴火の提起〜中部九州に分布する山甲川流紋岩を例として〜」
- 5月 15日 高木憲明 「地震計アレイを用いた阿蘇山・火山性連続微動の震源決定」
- 5月 21日 北田直人「重力異常からみた大阪盆地の地下構造」
- 8月 1日 宮澤理稔「受動的構造および能動的構造における地震波散乱問題」
- 11月 7日 岩村公太氏「流動不安定による火山性微動の励起シミュレーション」

2003年

2月 26日 宮澤理稔「1999年集集地震の遠地地震波到達直後に観測された、頻発する阿蘇火山性微動」

学 会 活 動 Activities in Scientific Societies

橋本武志

運営委員:地球電磁気·地球惑星圏学会

大沢信二

国際交流・広報委員:日本温泉科学会

編集委員:日本地熱学会

須藤靖明

編集委員:日本地震学会

竹村恵二

評議員:日本第四紀学会 編集幹事:第四紀研究

由佐悠紀

評議員:日本温泉科学会・日本地熱学会・日本陸水学会

運営委員長:陸水物理研究会

社会活動 Public Relations

大沢信二

大分県温泉調査研究会理事

新エネルギー・産業技術総合開発機構:地熱開発促進調査委員会(白水越地域部会委員

小野博尉

阿蘇町町史編集委員

須藤靖明

総合学習研究会(熊本県内小中学校先生と2ヶ月1回)

桜島運営協議会委員

火山噴火予知連絡会委員

阿蘇火山ガス安全対策専門委員会委員

くじゅう山系(硫黄山)防災協議会委員

九重山系火山砂防基本検討委員会委員

竹村恵二

文部科学省 科学技術政策研究所 科学技術動向研究センター 専門調査員

関西国際空港 (二期地区) 地盤挙動調査委員会委員

京都府地域活断層調查委員会委員

地震調查研究推進本部地震調查委員会北日本活断層分科会委員

京都盆地地下構造調査委員会委員

大分県地域活断層調査委員会委員

三重県地域活断層調査委員会委員

三重県伊勢平野地下構造調査委員会委員

断層研究資料センター第 15 回セミナー、一近畿地方土木地質図 改訂版の意義 「最近の後期新生代層の研究動向と土木地質」(2002 年 6 月 10 日、建設交流館グリーンホール)

宇治防災を考える市民の会講演、「京都盆地の地下構造と活断層-最近の調査結果」(2002年7月

19日、宇治職員会館)

ふるさと地盤診断ウオーク講師、地盤工学会関西支部(2002年10月12日)

オープンハウス地球熱学研究施設講演、「大分の活断層」(2002 年 11 月 9 日、ニューライフプラザ)ベイエリアの地盤情報と建設技術講習会、「大阪湾ベイエリアの地盤と自然環境一大阪湾の成り立ち」(2002 年 12 月 11 日、建設交流館グリーンホール)

地震セミナー、「京都盆地の活断層と地下構造」、京都市・文部科学省(2003年1月30日京都会館 大ホール)

亀岡断層調査現地説明会講師、京都府、(2003年2月14日)

田中良和 測地学審議会火山部会委員

火山噴火予知研究協議会委員

火山噴火予知研究委員会委員

地磁気観測作業委員会委員

火山研究センター一般公開特別講演講師

熊本県地質調査業協会講演会講師

電磁アクロスセミナー(2002年9月9日-10日、火山研究センター)

火山爆発の発生場と発生過程の研究会、火山研究センター(東北大、東工大、九州 大、京大防災研)

Hernandez A. Pedro: 中部アメリカにおける地震・火山災害の軽減 (ITER) (2002 年 11 月 27 日) 由佐悠紀

経済産業省原子力安全·保安院環境審查顧問

NEDO 地熱開発促進調查委員会委員

大分県環境審議会委員

大分県自然環境保全審議会委員 (温泉部会長)

大分県温泉地保全検討委員会委員長

大分県温泉監視調査委員会委員長

大分県天ケ瀬温泉保護対策検討委員会委員長

大分県土地収用事業認定審議会委員

大分県温泉調査研究会会長

日本温泉協会学術部委員

大分県原子爆弾被爆者対策協議会理事

別府市誌執筆者

別府市ふれあい広場・サザンクロス講師(2002年6月)

歴史と自然を学ぶ会十曜講座講師(2002年5月)

第16回九州都市教育長協議会定期総会特別講演講師(2002年10月)

京都大学地球熱学研究施設一般公開特別講演講師(2002年11月)

温泉療法フォーラム in 別府 パネラー (2003年1月)

別府温泉語り部の会講師(2003年2月)

京都大学地球物理学教室同窓の集い講演会講師(2003年3月)

オープンハウス報告

【別府】

京都大学大学院理学研究科附属地球熱学研究施設(別府)では 平成14年11月9日(土曜日)にオープンハウスを朝9時半から午後4時まで行いました。また、特別講演会と称し、午後2時と3時から、隣接するニューライフプラザの2階視聴覚室で2つの講演会を行いました。施設内では、10種類の公開実験または研究紹介展示を用意しました。合計205名の市民のみなさまにご来場いただきました。さらに、120名の方にアンケートで感想やご提案をお聞きすることができました。運営の方法、今後の運営に関する反省点とともに、オープンハウス当日の講演、展示、公開実験の概略、アンケートの集計結果をまとめたパンフレットを作成しました。昨年同様、教育委員会、市役所記者クラブ、新聞各紙、テレビ各局、ラジオ局のご協力を頂きました。ここに御礼申し上げます。以下に特別講演の題名と見学コースの名称と担当者を記します。

特別講演

竹村:「大分の活断層」、由佐:「別府温泉の話」

公開実験見学コース

10種類の実験、展示を行いました。各会場に説明者が待機し、参加者は地図を参考に自由に見学しました。

別府地域の岩石(2階セミナー室、竹村、宮崎)

地球は石でできている(2階セミナー室、柴田、芳川)

阿蘇の湯だまりの色の話(2階セミナー室、大沢)

阿蘇火山大研究コーナー(2階セミナー室、北田)

クリーンルーム(1階クリーンルーム室、ボグダン、米沢)

質量分析の道具(地下質量分析室、チャンチン、本多)

マグマを顕微鏡でみる(地階電子顕微鏡室、西村、高橋、川本)

電子顕微鏡で見てみよう(地階電子顕微鏡室、西村、高橋、川本)

地震計を見よう(地下地震計室、由佐、山田、長谷(阿蘇)、北田(阿蘇))

間欠泉 (建物の前。大沢)

マグマを高温で観察する(裏のコンテナハウス、鈴木、阿部、鄭)

2002 年度オープンハウス責任者 川本竜彦、柴田知之

【阿蘇】

1. 目的

一般市民,特に地域住民・関係機関に,当センターの活動内容を広く知ってもらうことで,センターに対する関心・理解を得る.また,社会への学術的知識の還元・啓蒙をはかる.

2. 開催日時

平成14年10月26日(土) 9:30~16:00

- 3. 内容
- ポスター展示(約30点)による研究内容の紹介・火山学の一般向け解説
- 公開実験
- ・「地震計の稼動」(大倉)
- ・「磁力計による岩石の磁化測定」(橋本)
- ·「火山噴火実験」(吉川慎・長谷)
- ・「火山工作」(吉川美)
- ・「中岳第1火口で採取した湖水のPh測定」(大沢)
- 施設備品展示(旧型地震計等各種観測装置の展示・解説)
- 火山に関するビデオの上映
- 火山に関する書籍の閲覧供与
- パソコンによる展示
- ・「九州の地震活動リアルタイムモニター」
- ・「火山と地震に関する教育ソフト」
- 特別講演(午前:田中良和教授,午後:須藤靖明助教授 各1時間)
- 見学者パンフレット (大人用,子供用)を配布
- 阿蘇・九重火山の絵はがきを記念品として配布
- 4. 社会告知の方法
- A4・A3版ポスターを作成

配布先:赤水郵便局・アゼリア・阿蘇駅・阿蘇火山防災協議会阿蘇火山博物館・阿蘇青年の家・阿蘇山 測候所・一の宮町役場・井出酒店・ウィナス・株式会社キンキ・河陽郵便局・九州東海大学・熊本 大学・グリーンストック・セブンイレブン・たわら屋・地球熱学研究施設・地球物理学教室・長陽 村観光協会・長陽村教育委員会・長陽村役場・東工大草津白根火山観測所・ニコニコ堂・白水村役 場・花阿蘇美・マグマ食堂・民宿阿蘇の湯・ルナ天文台(50音順)

- 火山研究センターホームページによる公示
- マスコミによる報道
- 事前報道:每日新聞
- ・事後報道:NHKテレビローカルニュース
- 陽の長い一日の村美術館(長陽村秋季イベント総合パンフレット)

5. 見学者に関する集計

来場者:55名

住所	人	割合	年代	人	割合	どう知ったか?	人	割合
熊本市	14	36.8%	10代	1	2. 7%	長陽村パンフレット	18	56.3%
長陽村	12	31.6%	20代	9	24.3%	友人・知人から	10	31.3%
阿蘇町	2	5.3%	30代	4	10.8%	インターネット	3	9.3%
益城町	2	5.3%	40代	8	21.6%	雑誌・新聞	1	3.1%
一の宮町	1	2.6%	50代	6	16. 2%			
別府市	2	5.3%	60代	9	24.3%			
富合町	1	2.6%						
その他	1	2.6%						
不明	3	7.9%						
回答数	38			37			32	

面白かった展示・実験・見学ツアー	人	割合
火山噴火・火山工作(実験)	6	28.6%
中岳第1火口で採取した湖水のPh測定(実験・ポスター)	4	19.0%
地震計の稼動(実験)	2	9.5%
横穴(見学ツアー)	2	9.5%
磁場関係(ポスター)	1	4.8%
昔の機材 (ウィヘルト地震計等)	1	4.8%
中岳の赤熱現象(ポスター)	1	4.8%
日本の火山(ポスター)	1	4.8%
全部	3	14.2%
回答数	21	

6. まとめ

今年度は、長陽村のイベントにあわせて10月に開催した。前年度より来場者数が減少しているが、その原因として構造探査(北海道駒ケ岳)および集中観測(富士山)などが一般見学会直前にあり、社会告知時期が前年度より遅れた事があげられる。ただ、長陽村パンフレットに限っては、長陽村の発行時期にあわせたため、相対的に一般の方々の目に多く触れた結果が上の表にあらわれたと考えられる。次回は、もっと早い時期からの社会告知が重要である。さらに、社会告知用のポスターやホームページ・パンフレット等には、より具体的な内容およびプログラムなどを記載することが必要である。

展示に関しては、ポスターを貼り付けるボードの変更と立掛け式から貼付け式にしたため、前回よりも見やすく、説明するスペースも広く取れるようになった。しかし、我々にとってポスターはセンターの業務をわかりやすく説明するための便利かつ重要な手段であるが、所詮紙であるためリアリティが

無く,一般の方々にとってはスタッフの説明をイメージしにくいのではないかと思う.次回は,ポスター前に実際に使用した機材などを並べておく事が可能ならば,それは改善できると考えている.

公開実験および工作に関しては、大人向けの講義中、主に10代向け実験・工作として準備していたが、 センターへのアクセス方法が車に限定され、10代の参加人数が少数であったため、午前の講義前と午 後講義後に大人も一緒に参加していただいた。噴火実験は、その時間帯に来場してくださった方ほとん どが参加され、多少の失敗はあったものの上の表から察するに一定の評価をいただいたと思う。また、 親子や大学生に参加していただいた火山工作も同様に評価をいただいた。次回は、もう少し準備期間を 設けて、より完成度の高い実験・工作を準備したい。

次回に向けて,より多くの方がセンターを訪れやすくまたは訪れてみたいと思えるように,もう少 し早い段階から企画を練ることが重要であると考える.

来訪者 Visitors

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	/11	17151	

平成 14 年

4月2日 大分県生活環境部:安部裕部長・岩本隆二次長

4月4日 大分県生活環境部:渡辺課長補佐・渕主任;国際航業:向山 栄・佐野正明

5月8日 別府市議会議員:永井 正

5月16日 西日本技術開発:島田・藤井

5月20日 中央開発:鹿・細谷・徳丸

5月21日 国際地震工学研修(地震学コース)研修監理員:中野真紀子・Pornpen PHUNTUMAT (タ

イ)・AL-SABRI Mohammed Abbas Naji Ahmed (イエメン)

5月22日 応用地質九州:松山尚典

5月28日 野津原町:小野助役・一万田商工課長

5月29日 別府市:安部温泉課長ら2名

5月30日 理学部用度掛:中道秀久(物品検査)

6月2日-4日 京都大学理学研究科・地球物理:金田平太郎・中村洋介・杉戸信彦・谷 美由起・井

上 勉・土志田正二・松岡 暁・今村朋裕;東大・地震研:木村治夫(京都大学理学

研究科応用地球科学現地講義)

6月4日 青山小学校:小田教諭・3年生15人

6月7日 JAMSTEC: 巽研究領域長・佐藤

6月19日 ケビン・ジョンソン

6月21日 東工大:高橋栄一教授(特別セミナー)

7月1日 大分大学附属中学校:2年生3名

7月16日 西日本技術開発:松本・島田・長野;応用地質九州:松山尚典

7月17日 別府市民: 友永; 阪神コンサルタンツ: 横田・末廣

7月24日 別府市議会議員:岩男三男,公明党大分県本部事務長:森正則,別府市温泉課長・安

部和男

7月23日-25 大阪市大:日加三千宣;名古屋大環境学研究科:土屋理恵

8月3日 武蔵中1年:野本ふみひこ

8月5日 別府市議会事務局: 荒金則明次長

8月5日-7日 Univ. Colorado: Eric Canon; 地域・地盤・環境: 北田奈緒子

8月6日 西日本技術開発:藤井・島田・長野

8月7日-10日 京都大学理学部:北島・池田・吉田・岩下・中村(地熱学実習B受講生)

8月8日 別府市教育委員会:安部課長・豊永係長

8月12日 衆議院議員:江田康幸;別府市議会議員:岩男三男;公明党大分県本部事務長:森正

則

8月16日 九州地熱株式会社:田村社長・宮澤専務

8月22日-27日 東邦大学:高松信樹・他3名

8月31日 九州電力: 天本敬吾

9月2日-8日 京大人間環境:山崎恭子

9月11日 画家:財津秀邦;京大人間環境:古川邦之(宇野(京大人間環境)

9月12日 大分大学:千田 昇;大分県:後藤;応用地質九州:松山尚典ほか1名

9月22日 奈良教育大学:長友恒人ほか3名

9月28日 京大地球物理:余田成男

10月5日 (株)キッズカンパニー:清水宏行ほか3名(建物写真撮影)

10月8日 川崎地質:安間 恵

10月9日 大分大学:川野田実夫教授,学生4名

10 月 11 日 JAMSTEC: 巽研究領域長

10月24日 京大施設部:大塚正人,林宜男,速水徹,河村隆司

10月28日 鬼山地獄:小玉

10月29日 京大名誉教授:西村進

11月5日 JAMSTEC:木下理事, 巽研究領域長

11月6日 大分モンゴル親善協会:岡洋一郎会長

11月18日 荒金正憲別府大名誉教授,九重町:熊谷公生

11月27日 経理部:中川権治,村上春茂,田中正人(施設調査)

12月2日 神奈川県立田奈高校長:長谷川昇,同金井高校長:濱野厚ほか2名

12月4日 湯原浩三九大名誉教授

12月16日 雑誌(じゃらん)記者:平山昭子,国立水俣病総合研究センター:松山明人

12月16日 1月11日 東邦大学 恩田裕二

12月17日 大分県生活環境課:渕祐一

12 月 18 日 JAMSTEC: 巽領域長(20 日まで)

12月20日 アジア太平洋大学:学生2人

平成 15 年

- 1月9日 大分モンゴル親善協会:岡洋一郎会長、西技:島田・長野・藤井
- 1月10日 別府市温泉課:安部課長
- 1月15日 中国温泉視察団一行(7名),大分県国際交流センター(2名)
- 1月17日 (社)全国ビルメンテナンス協会撮影(2名)
- 1月27日 大分大:千田 昇;大分県消防防災:後藤;応用地質九州
- 1月31日 日本地研(株): 冷川,田口,片平: 松山尚典;京大防災研: 三村 衛;地域・地盤・ 環境: 山本浩司
- 2月15日·16日 長崎大学:松岡数充;奈良教育大学:西田史朗;同志社大学:林田明;京都造形芸術 大:原田憲一
- 2月17日 東京大学:小屋口剛博助教授(集中講義)
- 2月18日 韓国文化放送 (MBS) 取材班一行 (3名)
- 2月24日 京都大学学術情報メディアセンター: 櫻井ほか2名
- 2月25日 施設部:清水晶一,新井弘,福山啓三
- 3月5日 大分県生活環境課:渕祐一
- 3月10日 NHK: 木村和人, 理学部: 永田幸司, 園畑一幸(財産調査)
- 3月13日 大分県土木建築部:橋本則敏ほか5名
- 3月20日 理学部:植村博樹,近藤芳弘,野田航多
- 3月24日 京大化学研究所:長崎順一ら2名
- 3月26日 京大地球物理:福田洋一助教授, 東敏博助手, Dendi Surya Kusuma (インドネシア), 熊本大学:長谷中利昭教授,京大地球物理:青木,小松ほか4名, NHK 取材班一行 (3名)

【阿蘇】

平成 14 年

- 4月11日・19日 京都大学大学院 人間・環境学研究科院生 古川邦之
- 4月19日 東大生産技術研究所 藤森教授 他多数見学
- 5月9日 森林総合研究所 宮縁育夫氏 研究打ち合わせ
- 5月14日・16日 東工大 金嶋聡・高木憲明
- 5月14日 火山研ボーリングコア検討会 多数来研
- 5月22日 国際地震工学研修(地震学コース)研修監理員:中野真紀子・Pornpen PHUNTUMAT (タイ)・AL-SABRI Mohammed Abbas Naji Ahmed (イエメン)
- 6月15日 Catharine Hurst
- 6月20日 九州大学工学部 福岡氏 セミナー
- 7月20日 福岡・小沢(九州大学資源)地磁気観測の研究
- 7月23日 石井美樹・豊留修一(気象庁地磁気観測所)観測打ち合わせ
- 7月23日 熊本県地質調査業協会 4名、研究会の報告
- 7月31日·8月4日京大理 久家慶子·宮澤理稔·本多茂喜 課題演習 D2 実習生 6名
- 8月1日-4日 地球電磁気学実習生

- 8月4日-7日 地熱学野外実習生
- 8月31日 城森明(ネオサイエンス) 研究情報交換
- 9月3日-4日 京都大学大学院 人間・環境学研究科院生 斎藤武志 東工大院生 高木憲明
- 9月3日 総合人間学部 斉藤氏, 東工大 高木氏
- 9月9日・10日 熊沢峰夫ほか3名 (核燃料サイクル機構)、電磁アクロス研究会
- 9月9日・10日 核燃料サイクル機構 熊澤氏・長尾氏・中島氏・松本氏 セミナー
- 9月17日 大久保綾子(防災研究所)研究打ち合わせ
- 9月20日 辻 保代ほか10名(東大出版会)研修
- 9月20日 東京大学出版会一行 見学
- 9月21日-22日 東大地震研 川勝均・山本希 東工大 金嶋聡
- 9月24日·25日 谷口、平林、清水、石原、井口、神田 特定領域研究A01班会議
- 9月24日 平林順一(東工大)・石原和弘・井口正人・神田径(SVO)谷口充宏(東北大)・清水
- 洋(九大) 科研費研究打ち合わせ会議
- 9月25日・25日 国際電子工業 井出氏・江浦氏 観測系メンテナンス
- 9月30日 日鉄鉱コンサルタント 横井氏 営業
- 10月10日 読売新聞社 宇佐美氏 取材
- 10月25日 京都大学 施設部 大塚正人・速水徹
- 10月26日 一般見学会で約80名来研
- 11月2日・4日 鎌田浩毅ほか10名、阿蘇火山の地質巡検
- 11月5日-8日 京大総合人間学部 酒井敏他6名
- 11月5日-11日 東工大院生 高木憲明
- 11月5日·5日 高木·金嶋·岩村(東工大)微動観測、総合人間学部 酒井氏一行 赤熱観測
- 11月6日-11日 東工大 金嶋聡・岩村公太・米陀
- 11月8日 結城洋一(応用地質)空中磁気測量打ち合わせ
- 11月8日 応用地質(株)3名 営業
- 11月15日·17日 Malcolm Johnston (USGS) 研究打ち合わせ
- 11月15日・19日マルコム滞在、日米共同事業のデータ解析
- 11月26日 NEC 三栄・九州技研 営業
- 11月27日 Pedro A., 野津憲治、大野正夫 セミナー
- 12月5日 JAMSTEC 市來雅啓氏 見学
- 12月12日 事務官4名(経理部契約課課長補佐野田修平、支出掛西浦利行、主計法規掛藤原浩一、 主計第二監査掛坂本雄美

平成 15 年

- 1月7日 立野ダム工事事務所 北嶋清氏 気象データのコピー
- 1月27日 事務官3名(成田、堀田、鴨嶋) 技官面接
- 2月19日 東京大学 小屋口剛博氏 セミナー
- 2月25日 27日 東工大 金嶋聡・高木憲明 京大理院生 宮澤理稔
- 2月28日 阿蘇火山博物館 池辺伸一郎氏 研究打ち合わせ

